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#### **SCIENCE**

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# Geology of the cretaceous Paraná Magmatic Province at the central portion of Torres Syncline, Southern Brazil

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#### **ABSTRACT**

Geological maps of tropical areas demand special challenges due to weathering conditions, mainly where relief has low variability. Most part of Paraná Magmatic Province is included in this scenario. An exception to this situation is the southeastern portion of the province, in the state of Rio Grande do Sul, Brazil. The stratigraphic disposition of volcanic units and the characteristics of erosion of each one caused the exposition of four identified plateaus, in which outcrop, from bottom to top: (1) basalts from Torres/Vale do Sol formations; (2) Caxias do Sul dacite; (3) Barros Cassal sequence; (4) Santa Maria rhyolite. The contacts were traced using SRTM images (through altitude and slope maps), integrating previously published semi- to detailed maps and geochemical database from our group. The resulting map (1:250,000) contributes to the volcano-stratigraphic knowledge of this huge province, and allows us to suggest the review of silicic units, initially called Palmas Formation.

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#### **KEYWORDS**

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#### 1. Introduction

The Paraná Magmatic Province (PMP) is one of the greatest expressions of continental volcanism on Planet Earth, and preserves ca. 1,000,000 km<sup>2</sup> of volcanic rocks (Frank et al., 2009), which can commonly reach thicknesses greater than 1000 m (Almeida, 1987). About 90% of this province (in area) is located in Brazilian territory, occupying significant portions of the southern (Rio Grande do Sul, Santa Catarina and Paraná), southeastern (São Paulo and Minas Gerais) and center-western (Mato Grosso, Mato Grosso do Sul and Goiás) states. Tholeitic basalts predominate, which form laterally expressive (estimated volume of ~450,000 km<sup>3</sup> for extrusive rocks in Brazilian territory; Frank et al., 2009) but often discontinuous flows. They must have been, in large part, fed by dike systems, with central edifices notably rare (see Pacheco et al., 2018).

As in other Continental Flood Basalt provinces, the mapping of flow successions is specially challenging, given their petrographic and structural similarities. For this reason, geophysical tools are often used. Studies were successful using paleomagnetic secular variation (e.g. for Deccan, Chenet et al., 2009; for Paraná, Raposo et al., 2023), for example. In addition, complex denudation and erosive processes placed during- and after- the main magmatic event (see AFT ages, Gallagher & Brown, 1999; Hueck et al.,

2019 and references therein) were responsible for a significant varying loss of rock volumes throughout coast and country side of Brazil. In the case of the Paraná Magmatic Province, the intense weathering and the smooth relief of much of the occurrence area reflect the adverse conditions of elaborating a detailed geological map and, consequently, knowing the volcano-stratigraphy.

Since the first studies on the PMP in Brazil (e.g. Bellieni et al., 1986), the most detailed surveys have been carried out on its southern portion (e.g. Bellieni et al., 1986; Garland et al., 1995), where steeper reliefs favor continuous exposures of the volcanic package over extensive sections. Thus, the most detailed knowledge about the stratigraphy, geochronology and architecture of the PMP volcanism in Brazil is concentrated in this region. Even so, studies are based on geological profiles obtained along the main roads that cross the volcanic pile. Some publications of detailed and semi-detailed scale maps of specific regions have been published in recent years by the Geological Survey of Brazil (Horn et al., 2022), or as support for scientific investigations (Barreto et al., 2014; Guimarães et al., 2018; Janasi et al., 2011; Polo et al., 2018; Rossetti et al., 2018).

The volcanism in southern Brazil is characterized by the expressive occurrence of silicic rocks (dacites and rhyolites) of diversified compositions. They have been the focus of several studies given peculiar characteristics, and whose spatial and stratigraphic distribution was presented in a map of regional scale by Nardy et al. (2008). Not by chance, the record of this type of volcanism lies within a regional structure called Torres syncline (e.g. Waichel et al., 2012): a large depression (ca. 400 km wide), with gutter form, and whose axis dip to the interior of the continent (NW). Although there is still a lack of understanding how exactly regional structures are distributed within the basin, and up to which point they are controlling the distribution of magmatism, at least since Almeida (1987) these NW-SE axes (South to North: Rio Grande Arch, Torres Syncline, Ponta Grossa Arch and Goiania Flexure) are recognized as big tectono-magmatic lineaments associated to the sedimentary and magmatic deposition/erosion during Paleozoic and Mesozoic. They also influenced the posterior cenozoic historic of denudation and erosion (e.g. Cogné et al., 2011; Hueck et al., 2019; Karl et al., 2013).

Our initial studies in the Torres syncline focused on the silicic volcanism, and for that purpose were concentrated in the vicinities of Barros Cassal and Gramado Xavier towns (Rio Grande do Sul State), where a large number of outcrops were exposed by the construction of a highway. The first surveys revealed an important compositional variation, supported by petrographic and geochemical studies (Polo et al., 2018; Polo & Janasi, 2014a). Previously, regional maps only indicated the presence of rhyolites from the Santa Maria unit. By current chemical classification of silicic components of low-Ti sequence (Palmas type volcanism), Santa Maria corresponds to the SiO<sub>2</sub>-rich and TiO<sub>2</sub>-poor among six subtypes defined as Palmas Type (Garland et al., 1995).

Importantly, chemical variations are expressed stratigraphic limits, with pre-rhyolite sequences outcropping in valleys. The bottom of the valleys is occupied by dacites equivalent to the Caxias do Sul type (Garland et al., 1995). They are, in several sections, separated from the rhyolites by a succession of thin flows of more primitive composition (andesites to dacites) hitherto unrecognized in the literature, and here referred as Barros Cassal Sequence (Polo et al., 2018; Polo & Janasi, 2014b). These 'pre-rhyolite packages' (including Barros Cassal sequence) are absent in certain regions and locally result in a direct contact between Santa Maria rhyolite and the Caxias do Sul dacite. Southward, the rhyolite outcrops directly in contact with basalt from Vale do Sol Fm, with the absence of dacites. Thus, in order to understand the spatial distribution of volcanic units and their relationships, it was necessary to carry out a mapping on a semi-detail scale, which resulted in the geological map at 1:100,000 published by Polo and Janasi (2014a). Our later studies expanded into neighboring areas, aiming at understanding the distribution and compositional diversity of the Barros Cassal Sequence and the meaning of basic to intermediate volcanism in the province.

Volcanic rocks of different compositions and structures have different weathering characteristics (i.e resistance) due to surface physical-chemical processes, which in this case, might reflect the diverse contents of SiO<sub>2</sub>. The stratigraphic disposition at the area and the contrast of erosion patterns drew topographic 'plateaus', separated and highlighted by steep declivities. Systematic integration of chemical database, field surveys and semi-detail mapping (published by Polo & Janasi, 2014a; located on the W-SW of the area) allowed the elaboration of a geological map of ca. 15,500 km<sup>2</sup> area in the central-east portion of the Torres Syncline (Main Map).

#### 2. Methods

#### 2.1. Field surveys and chemical database

The geological map is based on systematic field surveys carried out by our group in successive campaigns between 2010 and 2019. The surveys were placed along the main access routes in the region, where good quality rock exposures are often found. Occasionally, outcrops in riverbeds and in active or abandoned quarries were studied. With the support from our geochemical database, soon to be published, the development of a remote technique based on the topographic signature was possible.

# 2.2. Geomorphology analysis - geoprocessing

SRTM (NASA Shuttle Radar Topography Mission) images were used to guide tracing contacts with three main products: 1. Elevation model; 2. Hypsometric map (intervals of 52 m; 15 classes) (Figure 1 (a)); 3. Slope map (Figure 1(b)). Discontinuities and faults are inferred based mainly on the altitude of occurrence and comparison with chemical data, within topographically expressive lineaments (e.g. when the same unit appears at different altitudes at each side of the same drainage).

Four plateaus were identified for the whole area (Figure 1(a)). They are separated by high slopes (Figure 1(b)). The lowest one (Figure 1, 1) comprises valleys and drainages in the northern part of the area, and gets progressively more expressive towards south. It marks the outcropping of basalts from Torres/Vale do Sol Fm. The second, third and fourth plateaus comprise Caxias dacite, Barros Cassal Sequence and Santa Maria rhyolite, respectively. It is important to note that units are not equally distributed, with discontinuities and hiatuses. The geological sections presented with the map also show the identified plateaus.

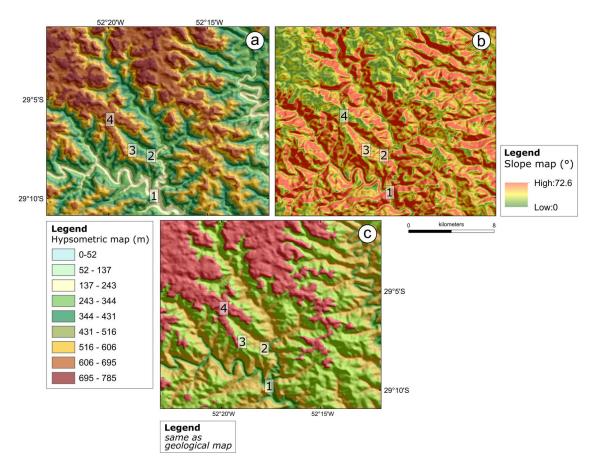


Figure 1. (a) Hypsometric map of classified altitudes, in intervals of 52 meters; (b) Slope map, that highlights 4 plateaus, indicated also in the geological sections; (c) Resulting geological map, after comparing geomorphologic patterns to geochemical and field data, besides integrating semi- to detailed maps within the area.

#### 3. Geological map - explanatory note

#### 3.1. Volcanic units

# 3.1.1. Torres and Vale do Sol formations

The mapped area is formed by volcanic rocks belonging to the Serra Geral Group and is dominated by silicic rocks. Earlier basalts occur only in the bottoms of the most carved valleys in its eastern part. In this portion, basalts from Torres and Vale do Sol formations (Figure 3(a-d)), that are dominant in terms of volume, occupy valleys and become dominant to the south of the area. Whenever observed, basalts unveil structural and geochemical characteristics compatible with the Vale do Sol Fm.; although it is possible that in deeper portions of the valleys the Fm. Torres may be present, no outcrops clearly associated with it have been identified. These two units are undivided in this map.

#### 3.1.2. Caxias do Sul dacite (CXd)

The Caxias do Sul-type dacite (CXd) (Figures 2 and 3 (e,f)) constitutes extensive deposits (apparently mostly lava flows and lava-domes; cf. Lima et al., 2012; Polo et al., 2018). The unit was presumably erupted continuously during the initial stage, whereas, as recorded by the deposition of sandstone layers interleaved with

the upper lava flows, a more intermittent eruptive behavior characterized the end of this event. Since it corresponds to the silicic rocks that occur over most of the area studied by Rossetti et al. (2018), and its stratigraphic position is consistent with their column, CXd is attributed here to the Palmas Fm.

In sections along the central and eastern part of the map, a layer of massive andesite is underlies the first CXd deposits (Figure 2); when observed, the contact is subtle, and was only detected after systematic gammaspectrometric measurements throughout the exposure, complemented by petrography and geochemistry (see below). The composition of the dacite in this outcrop is more primitive, akin to the Jacuí type (Garland et al., 1995; Nardy et al., 2008), which is a probable result of local interaction with the lower andesitic magma.

#### 3.1.3. Barros Cassal succession

The Barros Cassal succession is formed by the emplacement of multiple lava flows with frequent intercalations of thin sandstone layers and lenses (Figure 3(g)). It comprises a wide compositional range, from basalt to dacite, mostly with a well-defined trend of increasing silica content upwards; the more silicic rocks are more primitive than the CXd and occur exclusively

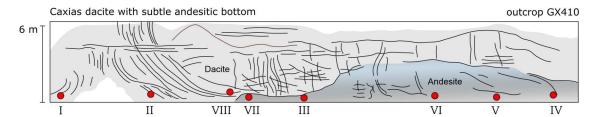


Figure 2. Sketch of outcrop GX-410 (UTM 385040 E, 6768730 N; 22J; Pouso Novo-Lajeado section) showing the relationships between CX dacite and underlying andesite. Red circles indicate location of sampling.



Figure 3. (a) Vesiculated basalt from Vale do Sol Fm., town of Herveiras; (b) Brecciated upper part of basalt from Vale do Sol Fm., town of Herveiras; (c) Pre-dacite basalt, Torres/Vale do Sol fms., West of Guaporé town; (d) Pahoehoe basalt of Torres Fm. with pipe vesicles; (e) Caxias dacite; (f) Vesiculated/amygdaloidal Caxias dacite; (q) Sandstone clastic dike within Barros Cassal dacite; (h) Santa Maria rhyolite; (i) Santa Maria rhyolite, town of Gramado Xavier; (j) Santa Maria rhyolite overlying directly pre-Caxias basalts, town of Herveiras.

in the western portion of the map (Gramado Xavier vicinities). The sequence changes eastward, as it tends to become more mafic (including basalts with up to 6 wt% MgO) and thicker (up to 200 m in the

Guaporé region), while it is thinned or even absent to the west in places where the dome structures of the previous Caxias dacite sustained a more prominent relief. The sedimentary intercalations disappear eastward, suggesting that the border of the basin was located to the west during the deposition.

#### 3.1.4. Santa Maria rhyolite (SMr)

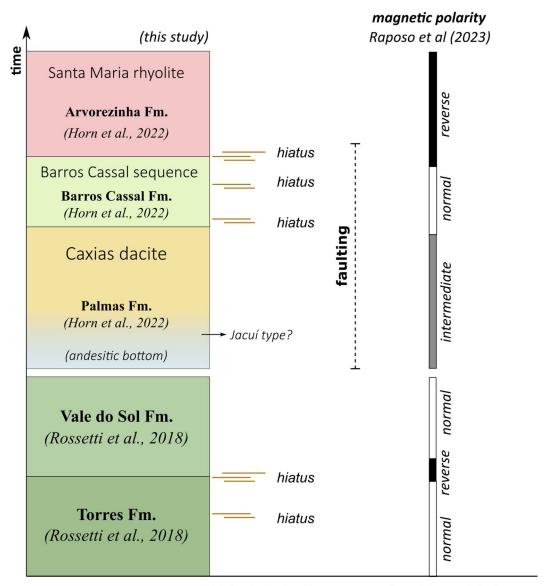
The upper volcanic sequence is the Santa Maria rhyolite (Figure 3(h-j)), an extensive and thick unit whose area of exposure is shown by regional maps (e.g. Nardy et al., 2008) to extend for at least 300 km to the west of our map along the axis of the Torres Syncline. In the mapped area, it may reach thicknesses up to 400 m, among the highest reported for this unit. In the western part of the map, the rhyolites usually rest directly on sandstones and volcanoclastic breccias (cf. Riccomini et al., 2016). Western of longitude 53°30′W (Nardy et al., 2008), Santa Maria rhyolite is found directly over the Vale do Sol Fm., as the previous Caxias Dacite and Barros Cassal Fm. fade out.

In most of the exposures the flow base is formed by up to 5 m thick layer of black obsidian. The lower deposits of the SMr have frequent structures indicative of effusive character, including lava lobes, autobreccias and peperite (produced by lava-sediment interaction) which are described in detail in Polo et al. (2018).

In terms of stratigraphy, Santa Maria rhyolite is here attributed to Arvorezinha Fm (from Horn et al., 2022).

### 3.2. Agglomerates, breccias and sandstones at the base of SMr

The sandstones are red colored, immature (feldsparrich), poorly selected, medium to coarse grained, with local conglomerate facies. They show some primary structures, such as plane-parallel layering and small to medium-scale cross-stratification (Figure 3 (g); clastic dike, sandstone cutting Barros Cassal unit). These structures and the presence of coarse facies and clasts (5-10 cm) dispersed in sandy matrix are suggestive of lacustrine-alluvial environments. Locally, these sandstones interacted with the rhyolite lavas of the Santa Maria forming peperites. An alternative explanation for the origin of some sedimentary



\*hiatus indicates the presence of sandstones, intertrapps, breccias, pepperites and/or conglomerates

**Figure 4.** Stratigraphy of the volcanic units from the Paraná Magmatic Province in the area of the map, referred to the stratigraphic columns adopted by Rossetti et al. (2018) and Horn et al. (2022) and to the paleomagnetic polarities obtained in the region by Raposo et al. (2023).

rocks in the southern Paraná as sand injectites and extrudites was proposed by Duarte and Hartmann (2014), but Polo and Janasi (2014b) argued for a syn-volcanic origin of the occurrences in the mapped area on the bases of field and mineralogical evidence.

#### 4. Discussions and final remarks

The most recent stratigraphic proposal for the Torres Syncline region (Rossetti et al., 2018) includes all the silicic volcanic rocks in a single stratigraphic unit, the Palmas Fm. It overlaps the initial basaltic sequences, represented by the Torres (pahoehoe basalts) and Vale do Sol (rubbly pahoehoe basalts and andesites) formations, and is succeeded by basalts from the Esmeralda Formation. Grouping of the entire silicic rock pack in the Palmas Fm. proved to be questionable. In fact, as

originally verified in the map by Polo and Janasi (2014b), the Caxias do Sul dacite and the Santa Maria rhyolite occupy different stratigraphic positions, and are separated over practically the entire area of the map by an expressive pack of more mafic rocks of the Barros Cassal sequence. We propose here a threefold division of the upper volcanic sequence of the Torres Syncline where the silicic rocks occur (Figure 4).

The Caxias do Sul dacite immediately overlaps the Vale do Sol Formation, and in the central part of the map area has an andesite layer at its base (*andesitic bottom*; Figures 2 and 4). A diffuse contact with this andesite and a more primitive character of the dacites that are immediately above it in the Pouso Novo section suggests some interaction between the two magmas. Moreover, the Jacuí-type dacites reported in the regional literature (e.g. Nardy et al., 2008) appear to occur in this same

stratigraphic position, and possibly has an origin similar to these 'contaminated' Caxias dacites.

The Santa Maria rhyolite is and independent stratigraphic unit, and is almost always separated from the Caxias dacite by a succession of mafic to intermediate lava flow, which is here proposed to comprise a separate unit, named Barros Cassal Formation. This unit shows important compositional variation from the west to the east portion of the map area: more evolved terms (dacites) cease to occur, basalts and andesites predominate and a layer of primitive basalts appears the basis of the sequence. These primitive basalts are chemically similar to the Esmeralda type basalt of Peate (1997), suggesting a possible correlation between the Barros Cassal Fm. and the Esmeralda Fm. of Rossetti et al. (2018), to be investigated in future surveys.

As demonstrated here and by others, volcanoclastic rocks (named Jacuí Group by Riccomini et al., 2016; see also Polo & Janasi, 2014a,b) and an up to 10 m thick layer of sandstone occur locally at the base of Santa Maria rhyolite. Thin (< 5 m thick) layers of vitrophyre are common, especially at the lower portions of the unit. A thick monotonous, compositionally homogeneous, platy granular rock which at least in part may have a pyroclastic origin (compare Luchetti et al., 2018; Polo et al., 2018). Paleomagnetic studies (Ernesto et al., 1990; Raposo et al., 2023; Figure 4) systematically show a reverse polarity for this unit.

Inferred faulting appears to be placed in between deposition/erosion of volcanic units, but before the deposition of the Santa Maria rhyolite. When mapped, units are subjected to abrupt discontinuities, changes in topography, *hiatus* indicators (intertrapps, breccias, conglomerates with chards and glass content). None indicative of faulting is observed in the field whatsoever, which could suggest an irregular filling pattern since pre-Caxias basic units.

#### **Software**

Geoprocessing was performed using ESRI ArcMap® software.

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No potential conflict of interest was reported by the author(s).

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#### Data availability statement

The authors confirm that the data supporting the findings of this study are available within the article and its supplementary materials.

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