

Arenigian tholeiitic basalts in the Famatina Ordovician basin, northwestern Argentina: emplacement conditions and their tectonic significance

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ABSTRACT. This study is focused on the analysis of volcanic deposits that crop out at the middle portion of the Las Planchadas range, northern part of the Famatina System in Argentina. These volcanic rocks are records of an Ordovician effusive basaltic volcanism that took place under subaqueous marine conditions. Along the study area crop out an Arenigian volcanic and volcanoclastic rocks succession with massive and autoclastic lavas, hyaloclastites of basaltic composition and volcanoclastic sandstones and mudstones. Large volumes of the volcanic deposits were strongly affected by fragmentation processes during their subaqueous emplacement and *in situ* accumulated as basaltic breccias. The same volcanic-volcanoclastic association crops out to the south of the Las Planchadas range, forming a basaltic belt with similar characteristics. The geochemical features of the basalts are compatible with depleted mid-ocean ridge basalt (MORB)-like source for the magma, with contribution of subducted related components such as water rich marine hemipelagic sediments, compatible with a back arc geotectonic setting developed along the northern part of the Famatina System during the Arenigian.

Keywords: Tholeiitic basalts, Syn-eruptive Hyaloclastic deposits, Ordovician, Famatina belt.

RESUMEN. Basaltos toleíticos arenigianos en la cuenca ordovícica famatiniana, noroeste de Argentina: condiciones de emplazamiento y su implicancia tectónica. El presente trabajo se enfoca en el análisis de los depósitos volcánicos que afloran en el tramo medio de la sierra de Las Planchadas, norte del Sistema de Famatina, Argentina. Estas rocas representan un registro completo del volcanismo ordovícico, correspondiente a efusiones basálticas emplazadas en un ambiente submarino. En la zona estudiada aflora una sucesión volcánica-volcanoclástica representada por lavas macizas y autoclásticas de composición basáltica, hialoclastitas basálticas, psamitas y pelitas volcanoclástica. Extensos volúmenes de estos depósitos fueron afectados fuertemente por procesos de fragmentación durante su emplazamiento subacuático, acumulándose *in situ* como brechas basálticas. Asociaciones volcánicas-volcanoclástica con idénticas características también afloran hacia el sur, en la misma sierra de Las Planchadas, y forman un cinturón volcánico de basaltos. Las características geoquímicas de estas volcanitas son comparables con basaltos empobrecidos, originados en zonas de dorsales oceánicas (MORB) y con componentes propios de subducción, como la participación de sedimentos hemipelágicos ricos en agua. Estas características son indicativas de un ambiente tectónico de retroarco ligado a la evolución del tramo norte del Sistema de Famatina durante el Arenigiano.

Palabras clave: Basaltos toleíticos, Depósitos hialoclásticos Sin-eruptivos, Ordovícico, Cinturón famatiniano.

1. Introduction

The studies on the ancient volcanic-volcaniclastic subaqueous rocks sequences carried out over the last decade, have been critical in reconstructing eruptive and depositional processes related to the evolution of many orogenic belts (*e.g.*, Brown *et al.*, 2002; Busby *et al.*, 2006; White, 2000). In the analysis of ancient successions of basaltic flows and associated volcaniclastic deposits emplaced in subaqueous environments, the works of White and Busby-Spera (1987), Templeton and Hanson (2003) and Caroff *et al.* (2009) among others, have provided substantial elements on the basis of field studies, analysis of structures and textures, for the understanding of the eruptive and depositional processes that controlled their formation. The study of Ordovician volcanic rock sequences along the Puna and the Famatina System (Northwestern Argentina) is the key to define their evolving tectonic setting at that time. In recent years these rocks sequences has been the subject of several studies (*e.g.*, Koukharsky *et al.*, 1996; Coira, 2008; Coira *et al.*, 1999, 2002, 2009; Viramonte *et al.*, 2005).

The Famatina System *sensu* Petersen and Leanza (1953) comprise a set of ranges extending north-south approximately between 27° and 31° S, along the La Rioja and Catamarca provinces. It extends between the Precordillera (northern portion)-Cordillera Frontal to the west and the Sierras Pampeanas to the east. The Famatina System is composed by a nearly 3,000 m thick rock sequence including Late Cambrian to Tremadocian carbonates and siliciclastic rocks, the Arenigian to Llanvirnian? volcanic-sedimentary deposits, and several Early Ordovician arc-related intrusives. These units have been studied by several authors (synthesis have been published in Aceñolaza *et al.*, 1996; Mángano and Buatois, 1996; Saavedra *et al.*, 1998; Astini, 1999; Mángano *et al.*, 2003; among others), suggesting different tectonic setting interpretations. At the northern portion of the Famatina System, along the Las Planchadas and Narváez ranges (Fig. 1), the most voluminous volcanic and volcaniclastic deposits of proved Arenigian age are exposed (Turner, 1967; Cisterna *et al.*, 2010a, b). The study of these volcanic rock sequences may be one of the keys to understand the evolution of the Famatinian terrains as a portion of the western margin of the Gondwana.

This research documents the lavas and syneruptive products of an Arenigian basaltic volcanism occurred in a subaqueous environment. It develops large and well exposed records along Las Planchadas range (north of the Famatina System, northwestern Argentina, Fig. 1). The study of these deposits was made pointing on (i) the existence of several well-preserved outcrops of basalts displaying characteristic features of an ancient volcanism emplaced in a submarine environment, including well-developed hyaloclastic facies and spatially related to the lavas. (ii) The possibility to relate the studied basalts with other volcanic deposits that crop out in the same geological context and display petrographical and chemical similarities. (iii) To interpret the geodynamic evolution of the Ordovician volcanic belts, at the north portion of the Famatina System as a piece of the Puna - Famatina Belt in the regional context of the Southern Central Andes.

2. Regional geotectonic interpretations and stratigraphic framework

The tectonic setting of the Famatina System during the Early Palaeozoic has been a matter of debate during the recent years. While petrologic studies consider this belt as being built on the western Gondwana margin, on the other hand paleomagnetic and biogeographical studies (Conti *et al.*, 1996; Astini and Benedetto, 1996; Benedetto, 1998; Rapalini *et al.*, 1999) point to a parautochthonous or allochthonous intra Iapetus island arc system. The island arc volcanic setting for the Famatina System (located in the central domain of the Famatina Belt) was initially suggested by Aceñolaza and Tosselli (1984). Subsequently, some authors have favored a back arc setting for that belt on petrological basis (Mannheim, 1993; Clemens, 1993; Toselli *et al.*, 1996). Others, analyzing the stratigraphic and sedimentological records, have indicated an intra-arc basin evolution (Mangano and Buatois, 1996; Cisterna *et al.*, 2010a). Isotopic and geochemical studies on Famatina Early Ordovician granitoids documented the evolution of that magmatic belt on a thickened continental crust (Pankhurst *et al.*, 2000), instead of a genesis under an island arc regime. The extension of this magmatic belt northward in the Puna magmatic belts (Fig. 1) has been postulated by Coira (1975, 1979), Coira *et al.* (1982), Ramos (1988), Conti *et al.* (1996) and Cisterna *et al.* (2010b).

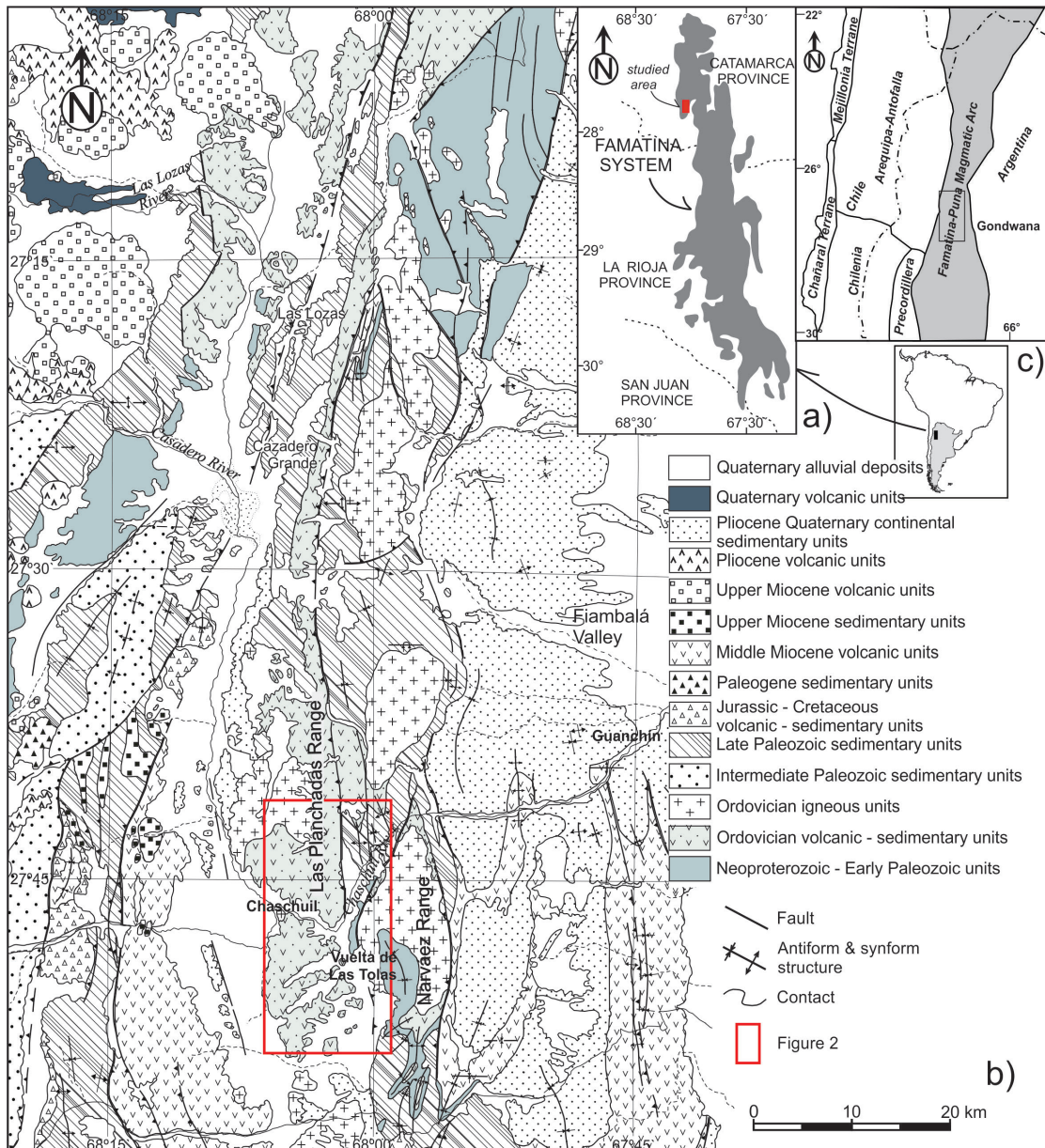


FIG. 1. **a.** Geological setting of the Famatina-Puna Belt; **b.** Geological map of the northern Famatina System; **c.** Regional context of the Famatina-Puna Magmatic Arc.

The Ordovician rocks in the northern portion of the Famatina System comprise Tremadocian and Arenigian volcanic-sedimentary successions. The Tremadocian rocks (Cisterna *et al.*, 2010b) are intruded by the Las Angosturas Granodiorite (Cisterna, 1994) with an U/Pb age of 485 ± 7 My (Rubiolo *et al.*, 2002). Nevertheless, the most extensive deposits

are represented by the Arenigian succession (Suri Formation, Harrington and Leanza, 1957; and Las Planchadas Formation, Turner, 1958) mainly exposed in the Las Planchadas range (Fig. 2).

To the south of the Las Planchadas range (Chaschuil area) the Arenigian succession is composed by dominant volcanoclastic lithofacies represented by

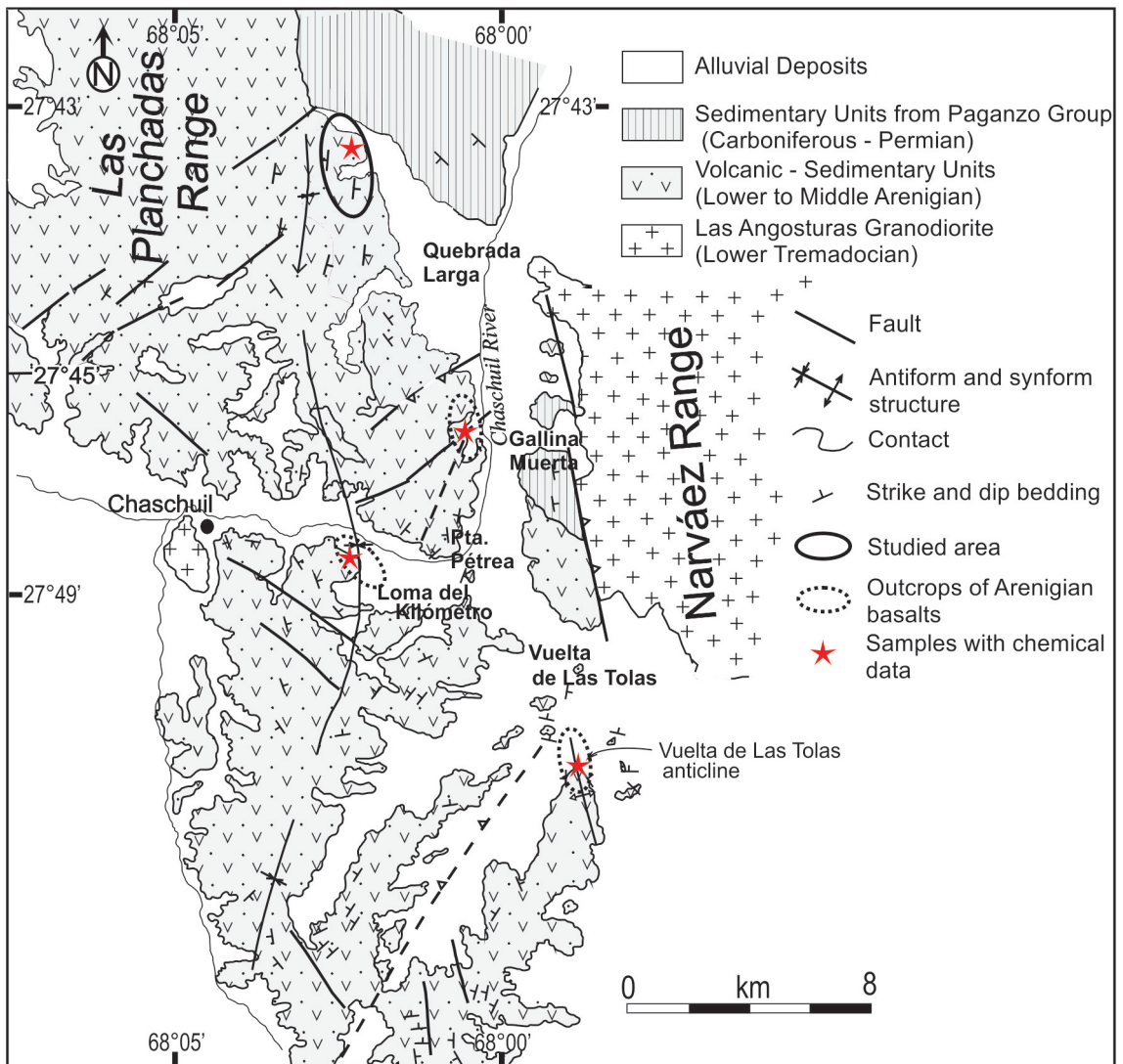


FIG. 2. Map of middle to southern portion of the Las Planchadas range and north of the Narv ez range.

volcaniclastic debris flows, turbidity currents and minor resedimented syn-eruptive pyroclastic deposits and lavas (Cisterna and Coira, 2014). At the upper portions, the rock sequence is represented by a volcanogenic sedimentary lithofacies with fossiliferous levels (M ngano and Buatois, 1997). The association of a brachiopod-rich fauna associated with trilobites and conodonts in the pelitic levels allows to assign this unit to the Arenigian (Albanessi and Vaccari, 1994; Benedetto, 1994; Vaccari and Waisfeld, 1994).

At the central portion of the Las Planchadas range (Quebrada Larga-de La Gallina), large outcrops of the Arenigian deposits are mainly composed by basaltic, andesitic and dacitic lavas, with significant volumes of lava-derived clastic deposits (Figs. 2 and 3), produced by quench fragmentation and gravitational collapse of lava deposits (Cisterna *et al.*, 2010a; Slavutsky, 2009). The upper levels of these deposits display volcaniclastic breccias, tuffs and lapilli beds, intercalated with volcanogenic fossiliferous facies, of Arenigian age (Cisterna *et*

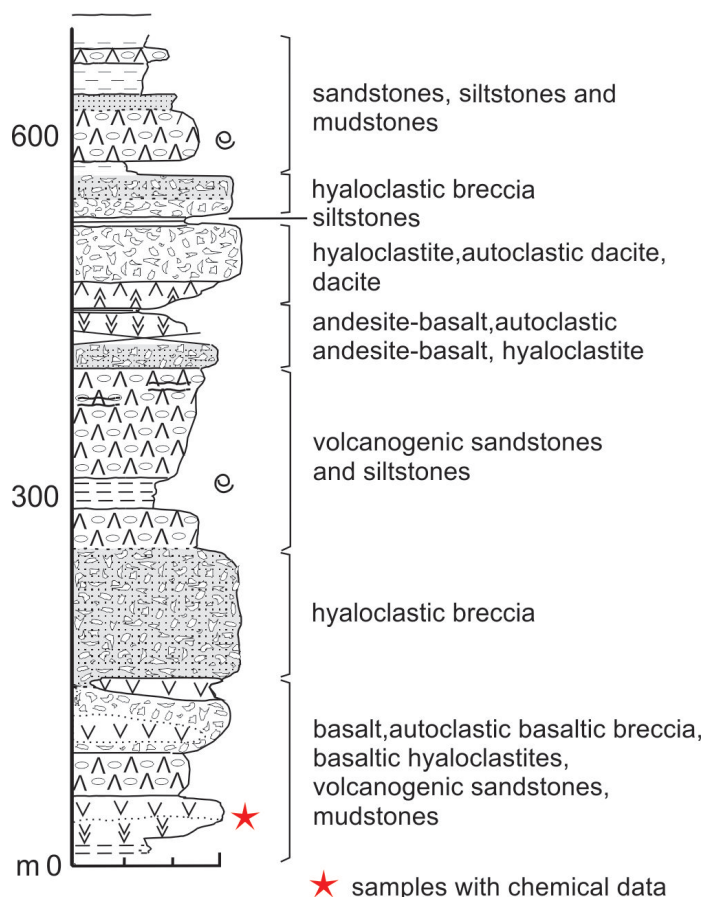


FIG. 3. Generalized stratigraphic section of the Arenigian volcanic deposits cropping out at the Las Planchadas range. The studied succession is part of the lower portion of this column.

al., 2010a). The studied succession represents the lower level of those deposits and the most basic ones.

The Ordovician volcanic successions are covered, in unconformity, by thick Late Paleozoic siliciclastic continental units (Turner, 1967). These sedimentary deposits were defined as the Agua Colorada and Patquía formations and were assigned to the Carboniferous and Permian, respectively (Turner, 1960; Cuerda, 1965).

3. Internal stratigraphy and lithofacies descriptions at the basaltic members

The present study was focused on well preserved volcanic deposits (nearly 5 km²) that crop out at the middle portion of the Las Planchadas range (Figs. 2 and 4a). They bring together the different facies that

characterize the Arenigian volcanic and volcanoclastic deposits along the Las Planchadas range (Quebrada de La Gallina, Vuelta de Las Tolas and Chaschuil, Figs. 2 and 5). This Ordovician stratigraphic succession builds up a synclinal structure, whose axis is subhorizontal, with a north-south trend. The field mapping shows that the Ordovician deposits certainly form a sequence without tectonic repetition. The base of the section is poorly exposed to the east, where locally an irregular contact can be observed between the basalt and a yellow siltstone (Fig. 4b). The top of the section has volcanoclastic facies (sandstones and mudstones), laterally equivalent to the fossiliferous volcanoclastic sandstones containing an Arenigian brachiopod fauna (Cisterna *et al.*, 2010a) cropping out along the Quebrada Larga-de La Gallina (to the south of the study area) (Figs. 2 and 3).

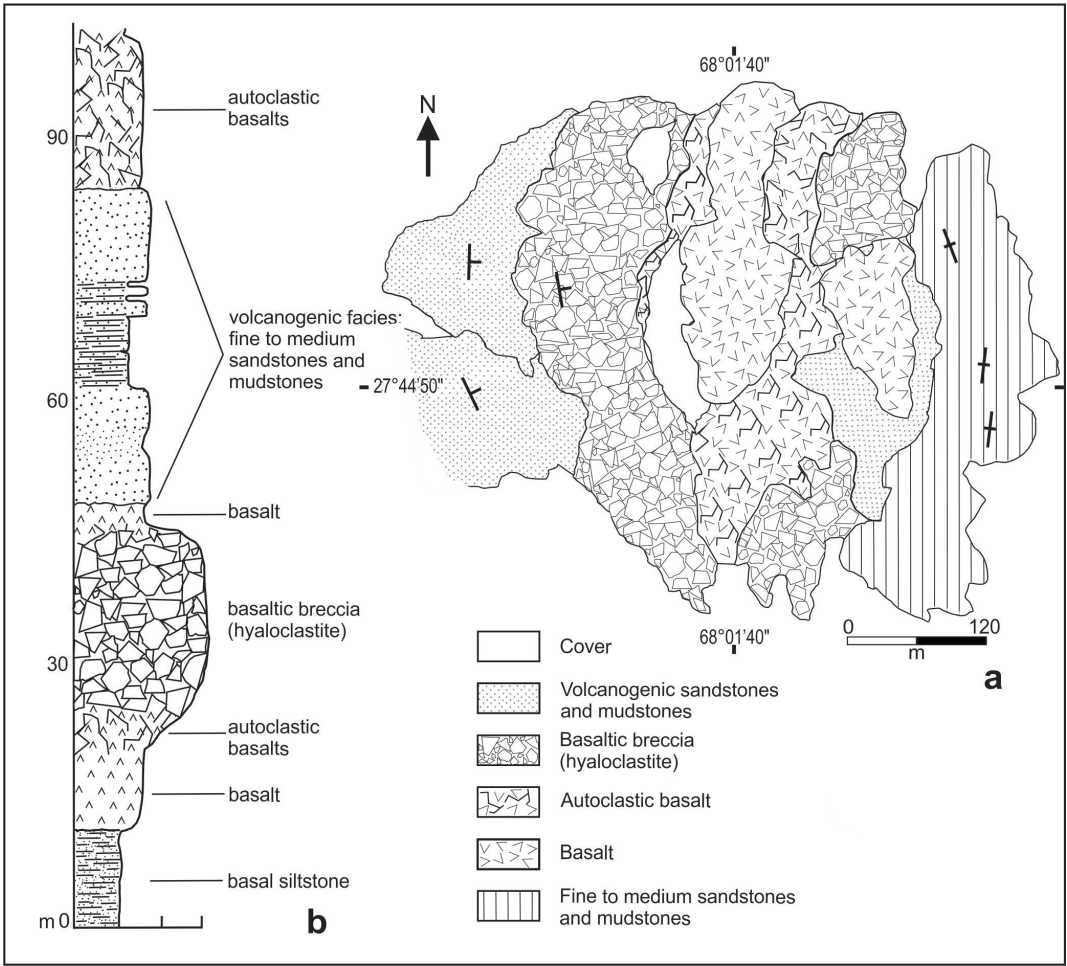


FIG. 4. a. Map of the basaltic volcanoclastic deposits studied; b. Schematic log showing the internal organization of the analyzed section.

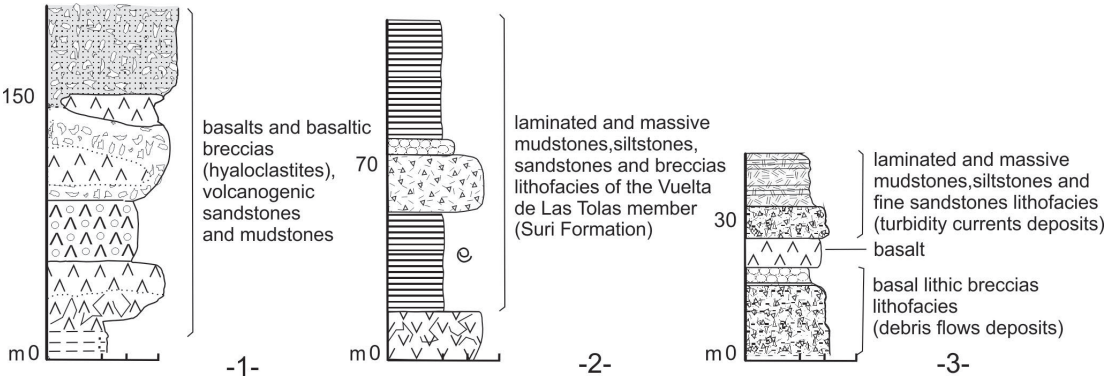


FIG. 5. Generalized comparative logs of the 1. Lower portion of the Quebrada Larga-La Gallina section (simplified from Cisterna *et al.*, 2010a); 2. Lower portion of the Vuelta de Las Tolas section; 3. Intermediate portion along the Chaschuil section (simplified from Cisterna and Coira, 2014).

Although the majority of the studied rocks have been affected by hydrothermal and low grade metamorphic processes, it is still possible to recognize their primary petrographic characteristics and to analyze their chemical signature on the basis of immobile elements.

3.1. Facies descriptions

3.1.1. Basaltic facies

Basalts crop out as lenticular or tabular bodies, varying between 2.5-10 m thick and usually showing transitional relationships to the autoclastic facies (Fig. 4b). They are greenish black, generally massive and fine-grained with a groundmass of varying texture and crystallinity. By sectors, basalts show a flow foliated structure and irregular vesicles (2 mm-2.5 cm).

Generally the basalts are porphyritic, sometimes glomerophytic, and phenocrysts (5-10%) are mostly plagioclase (1-3 mm) (27%) and lesser amounts augite (1mm) (1.5%). The phenocrysts may be fractured, especially the plagioclase which is corroded and frequently shows numerous glass inclusions (Figs. 6a, b). The scarce augite phenocrysts (Fig. 6d) are euhedral, while augite microphenocrysts are mostly equant, and may show a compositional zonation, almost always overgrown by tremolite (Figs. 6e and 7a). Olivine also forms microphenocrysts, frequently replaced by pyroxene and/or amphibole and may be fragmented (Figs. 6e, f), displaying variable degrees of alteration to iddingsite and carbonate. The groundmass is grey to greenish grey; its texture is generally pilotaxitic and rarely ophitic or subophitic (Fig. 6c). Besides glass, which is commonly present, the groundmass constituents are plagioclase microlites, augite, olivine and abundant opaque minerals. The microlites are weakly to moderately align; some are deflected around the phenocrysts and show morphologies such as swallow-tail terminations. The glass may be either fresh or altered to palagonite (Fig. 7b). Vesicles, which may be flow-aligned, are filled by calcite, quartz and/or chlorite. Some round vesicles are filled by quartz (Fig. 7c).

3.1.2. Autoclastic basaltic facies

They are composed by highly fractured basalts, developing jigsaw-fit textures (Fig. 8a). The blocky fragments are larger than 3 cm, most of them have curvilinear fracture surfaces; no chilled margins are

observed. The contacts may be sharp or gradational between these deposits and the coherent basalts.

3.1.3. Hyaloclastic basaltic facies

They form deposits up to 60 m thick, in transitional or sharp contact related to the autoclastic facies and massive basalts. The hyaloclastites within the basalts are commonly observed as tabular bodies less than 1 m in thickness (Fig. 8a). Hyaloclastites are poorly sorted monomictic basaltic breccias, with randomly oriented clasts. The shapes vary between flow foliated slabs with jagged ends, and ragged or blocky clasts with curvilinear surfaces (Fig. 8b). Besides, they may also display a jigsaw texture. The clasts may develop quenched margins and glassy rims. Their size varies in different sectors (from less than one mm up to 50 cm), usually larger near the massive lavas. The groundmass is composed of glassy particles and phenocryst fragments (plagioclase), carbonate and palagonite. The vesicles are irregular and may show flow alignment (Fig. 8b). The clasts are hyaloclastic basalts showing different structures, as described below (Figs. 9a, b and c). The most frequent have a vitric groundmass, sometimes altered to palagonite, with thin plagioclase laths in a pilotaxitic groundmass and elongated or subrounded vesicles, aligned parallel to the long axis of the clasts, and filled by carbonate or chlorite (Figs. 9b, c). The basaltic clasts may also be rounded and subrounded, with quenched margins (Figs. 9a, d, e and 8a) and a groundmass composed by glass or oxidized glass, feldspar microlites and vesicles filled by carbonate, chlorite and quartz; phenocrysts, sometimes fractured, are plagioclase and augite. Plagioclase may be euhedral or display large corroded crystals (Figs. 10b, c). They frequently show zonation and numerous glass inclusions (Fig. 10d). The augite and olivine (as microphenocrysts) are overgrown by amphiboles (Fig. 7a). Other types of clasts are rounded and highly vesiculated basalts (Fig. 9f).

4. Geochemistry

4.1. Methodology

All geochemical analyses were performed on selected samples, after close petrographic observations. The localities of each sample can be seen at the figure 2 and they correspond to basalts (MQL3,

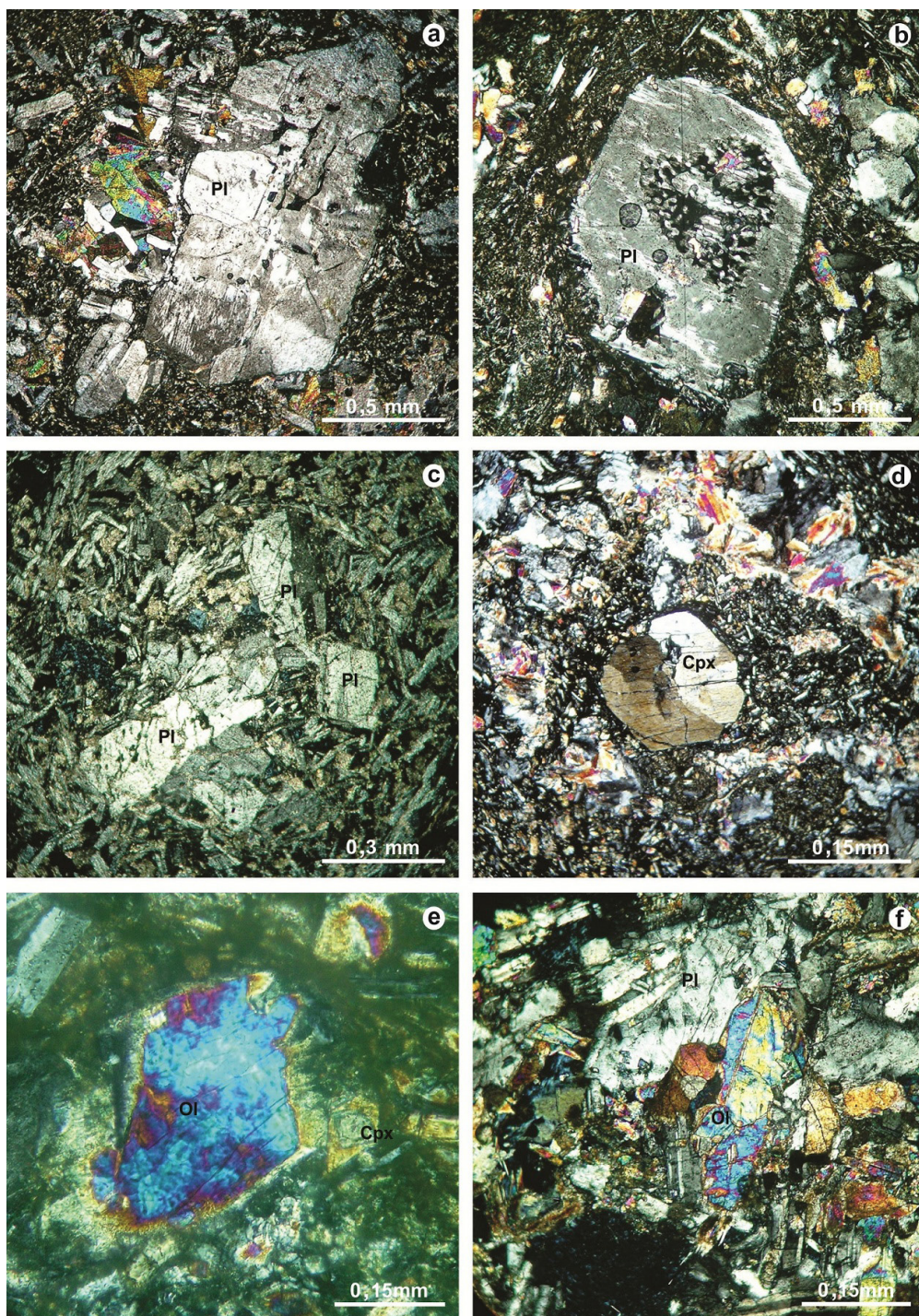


FIG. 6. Basalts showing: **a.** Plagioclase phenocryst showing resorbed zones near the core and fractures; **b.** Extensive glass inclusions in the core of the plagioclase phenocryst; **c.** Plagioclase phenocrysts developing glomerophytic clots and pilotaxitic groundmass; **d.** Equant augite microphenocryst, euhedral and zoned in a groundmass showing abundant devitrified glass; **e-f.** Olivine microphenocrysts partially resorbed and as fragmental grains. Crossed polarized light in all photomicrographs (Mineral abbreviations as suggested by Kretz, 1983).

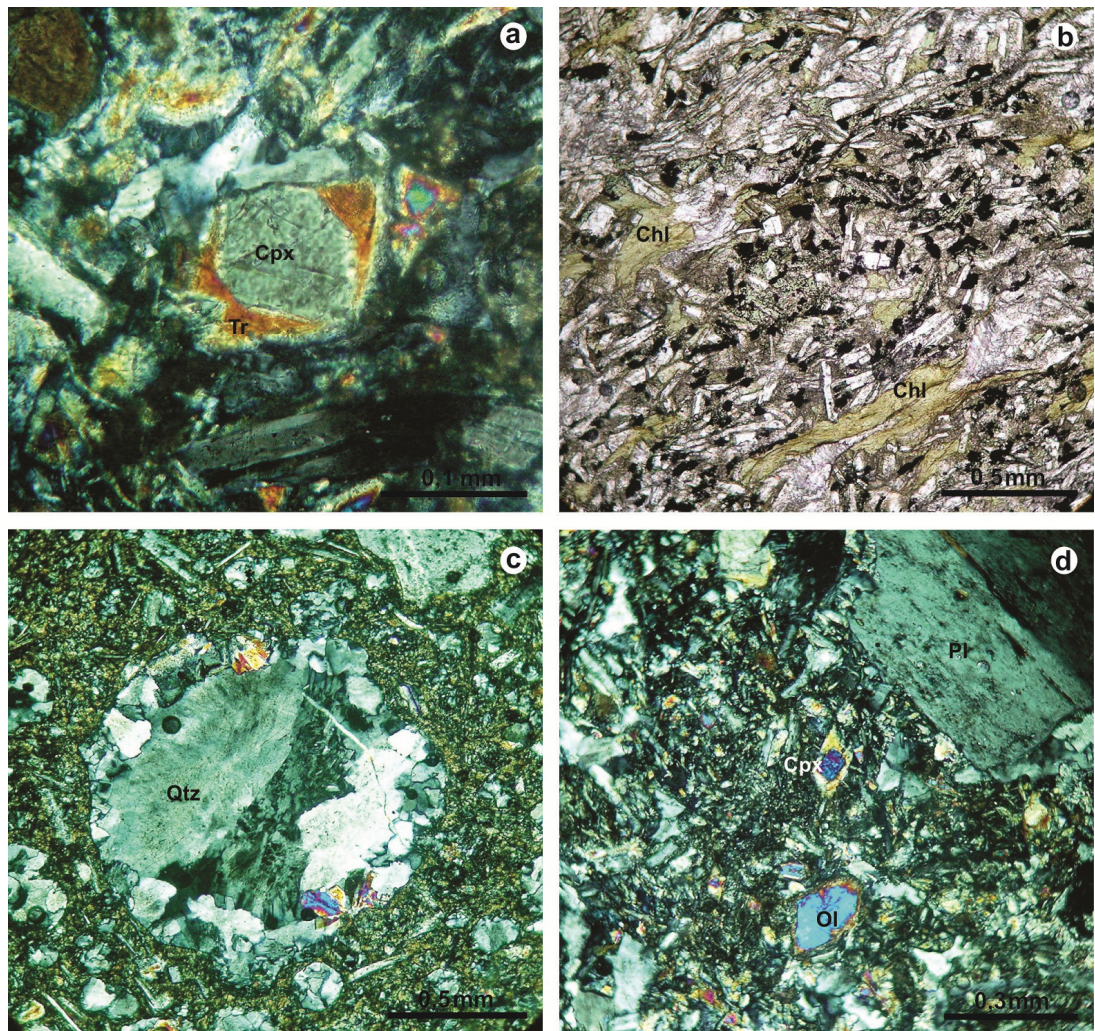


FIG. 7. Photomicrographs showing **a.** Augite overgrown by tremolite. Crossed polarized light. (Mineral abbreviations as suggested by Kretz, 1983); **b.** Plagioclase microliths, magnetite and palagonite replacing the glass groundmass. Plane polarized light; **c.** Rounded vesicle filled by quartz. Crossed polarized light; **d.** Basaltic groundmass showing, besides glass, plagioclase microlites, augite, olivine and abundant opaque minerals. Crossed polarized light.

MQL5, NQL1, NQL5, NQL18, NQL19, PC10, VT10, T5 and CG147) and basaltic hyaloclastites (NQL17, PC13, B11). The compositions of the whole-rock samples, including major, minor and trace elements were determined by wavelength dispersive X-ray fluorescence spectrometry (WD-XRFS). Rare earth element (REE) analyses were collected using a VARIAN Vista MPX ICP-OES (inductively coupled plasma-optical emission spectrometry). All the previous analyses were performed at either the GFZ in Potsdam or at the laboratories of the Institute of

Earth and Environmental Science, University of Potsdam, Germany.

The CG-147, PC10, VT10 and T5 samples were analyzed at the Chemical Laboratory of the Institute of Geosciences (Sao Paulo University) where major oxides and trace elements were determined by wavelength-dispersive X-ray fluorescence spectrometry (WD-XRFS) on fused glass discs using a Philips 9PW2400 spectrometer. Trace elements, including REE, were determined by inductively-coupled plasma mass spectrometry (ICP-MS).

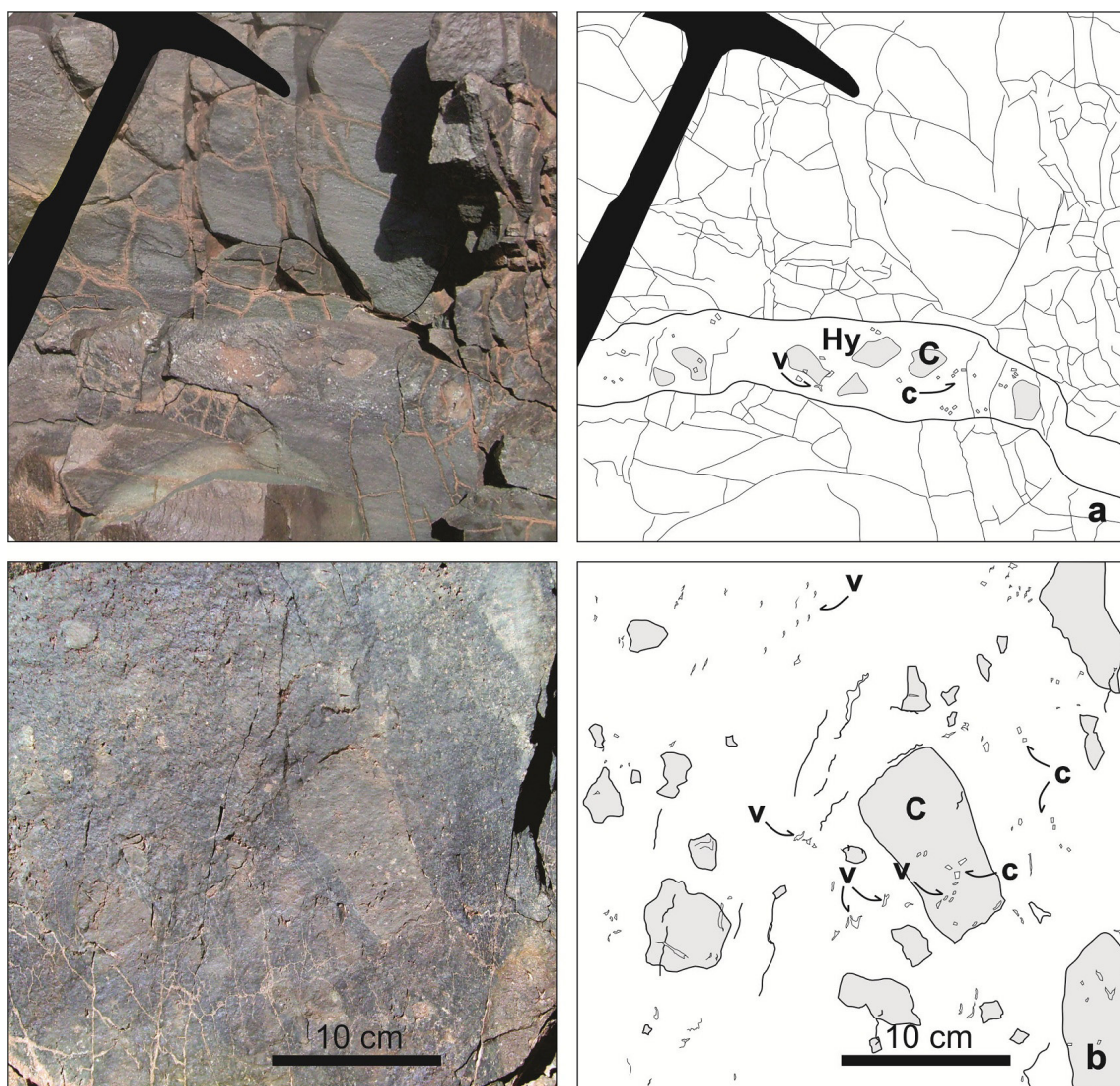


FIG. 8. **a.** Autoclastic basalt showing a jigsaw texture and a dyke of groundmass rich in *in situ* hyaloclastite. Clasts in the hyaloclastite have curvilinear margins and are oxidized. Vesicles are filled by carbonate. Schematic graph of the autoclastic basalt and the hyaloclastite (Hy). (C: clasts, c: crystalloclasts, v: vesicles); **b.** Monomictic hyaloclastite with clasts of different size and shape. They are blocky and the margins are not quenched. Clasts and groundmass contain abundant, very irregular vesicles. Schematic graph showing the monomictic hyaloclastite (C: clasts, c: crystalloclasts, v: vesicles).

4.2. Major and trace elements

Samples of basalts of the study area were analyzed alongside those of correlated basalts that crop out to the south (Quebrada Larga, Quebrada de la Gallina, Vuelta de Las Tolas and Chaschuil) (Fig. 2), in order to find out if they show compositional affinities that link them to the same magmatic event. The samples

include hyaloclastic basalts (from the Quebrada Larga) showing some variations regarding to the massive basalts, due to the presence of siliceous aggregates along the fractures. Results are shown in table 1.

Most of the samples show a relatively high LOI, so we have opted for their characterization by the use of immobile elements. For classification, the Zr/TiO₂ versus SiO₂ diagram of Winchester and Floyd

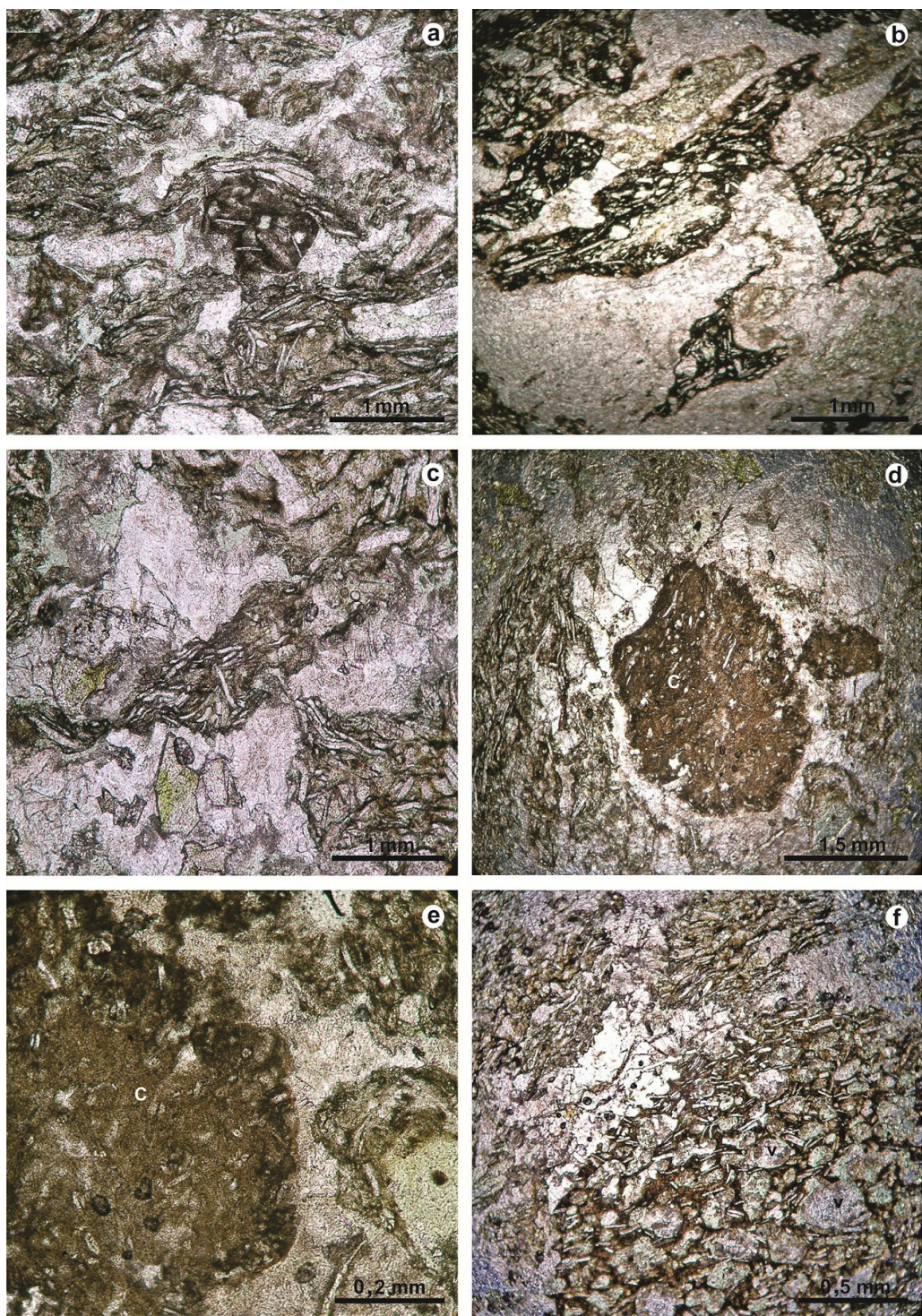


FIG. 9. Basaltic hyaloclastites showing at the microscope **a**. Small equant basaltic clast together with flow foliated clasts; **b-c**. Elongated clasts, generally with jagged ends or ragged. Calcite is abundant in the groundmass; **d**. Rounded basaltic clast with quenched margins, the detail (**e**) showing the oxidized glass groundmass; **f**. Rounded basaltic clast with abundant rounded vesicles filled by carbonate, chlorite and quartz. Plane polarized light in all photomicrographs.

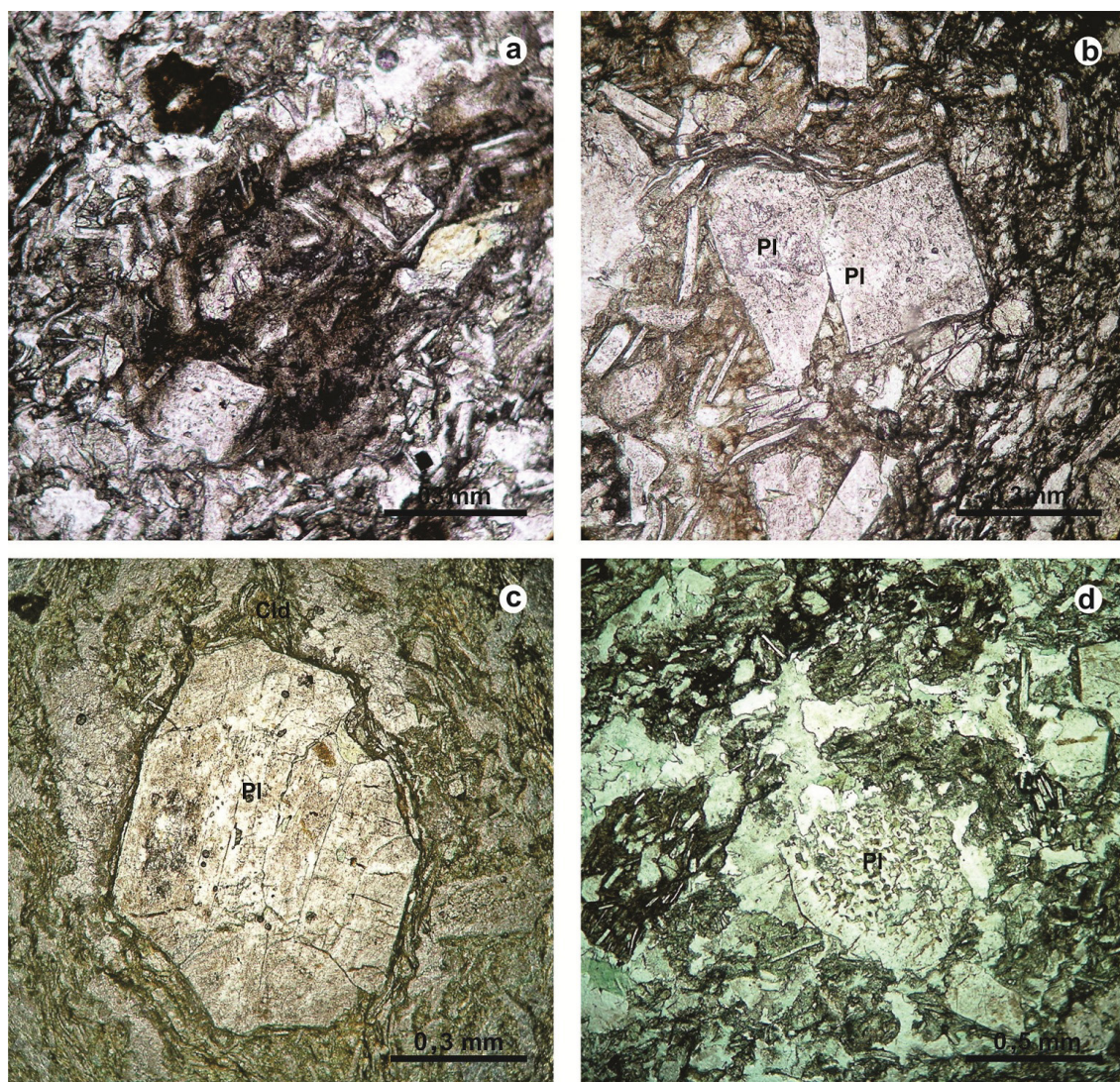


FIG. 10. Photomicrographs showing **a.** Small basaltic clast with glass groundmass, euhedral plagioclase microphenocryst and microliths; **b-c.** Euhedral and partially corroded crystalloclasts of plagioclase; **d.** Plagioclase crystalloclast showing glass inclusions in the core. Plane polarized light in all photomicrographs.

(1977) has been applied, indicating that they are geochemically basalts (Fig. 11a). The FeO^*/MgO versus SiO_2 diagram (Miyashiro, 1974) confirms the tholeiitic character of the basalts like the Zr/Y relation (Barrett and MacLean, 1999) (Figs. 11b, c). The hyaloclastic basalts are characterized by the lowest content of Fe_2O_3 , MgO and Al_2O_3 (Table 1) and the enrichment in SiO_2 (average 69.05% and 47.40% in massive basalts) that is related to the hydrothermal alteration, so these data are not used.

The analyzed basalts display low Mg-numbers [$\text{Mg\#} = 100 \times \text{Mg}/(\text{Mg} + \text{Fe}^{2+})$] ranging from 56 to 36 (Table 1), low Ni (20-63 ppm), Cr (85-199 ppm), MgO (3.94-7.45 wt%) and relatively high LOI (8.34-3.38wt%). Likewise they are low-Ti on the basis of their Ti/Y ratios=213-278 ($\text{Ti}/\text{Y} < 500$: low-Ti basalts, c.f. Cummins *et al.*, 1992) and $\text{TiO}_2 = 0.76\text{-}0.96\%$ (TiO_2 wt% < 2.5 : low-Ti basalts, c.f. Cummins *et al.*, 1992). They have an average of $\text{FeO}t/\text{TiO}_2 = 11.89$ and the trace elements

TABLE 1. MAJOR, TRACE AND REE ELEMENTS CONTENTS OF THE BASALTS AND BASALTIC HYALOCLASTITES FROM THE QUEBRADA LARGA AND OTHERS FROM CORRELATED BASALTS THAT CROP OUT TO THE SOUTH EASTERN AND SOUTH OF THE AREA.

Location	Northern Quebrada Larga							Southeastern Quebrada Larga		Quebrada de La Gallina	Vuelta de Las Tolas	Chaschuil	
Sample	NQL1	NQL5	NQL17	NQL18	NQL19	MQL3	MQL5	PC-10*	PC-13	B11	VT-10*	T-5*	CG-147*
Rock type	basalt	basalt	basaltic hyaloclastite	basalt	basalt	basalt	basalt	basalt	pillow basaltic breccia	basaltic hyaloclastite	basalt	basalt	basalt
Location	27°44'58.58" 68°01'37.85"	27°44'56.88" 68°01'38.93"	27°44'45.89" 68°01'47.42"	27°44'50.69" 68°01'44.18"	27°44'47.04" 68°01'48.41"	27°44'29.99" 68°01'38.84"	27°44'33.96" 68°01'41.30"	27°45'36.37" 68°01'50.11"	27°45'42.61" 68°01'42.07"	27°46'24.39" 68°01'38.37"	27°47'45.68" 68°01'44.90"	27°49'23.06" 68°02'33.84"	27°47'47.03" 68°03'36.01"
wt%													
SiO ₂	45.8	46.30	68.60	49.50	47.30	45.40	46.90	47.93	50.60	69.50	48.69	50.10	48.00
TiO ₂	0.92	0.96	0.25	0.78	0.93	0.81	0.78	0.96	0.83	0.20	0.90	0.76	0.79
Al ₂ O ₃	17.20	16.00	9.40	14.30	17.10	17.40	16.40	16.90	18.60	10.40	15.39	14.32	16.51
Fe ₂ O ₃	10.48	10.62	1.13	7.68	11.16	10.80	10.46	11.07	10.24	6.59	10.64	8.28	10.69
MnO	0.18	0.20	0.08	0.16	0.17	0.18	0.20	0.19	0.18	0.15	0.19	0.31	0.19
MgO	7.31	7.13	0.48	5.15	7.30	6.76	7.45	5.73	5.50	2.10	5.58	3.94	6.61
CaO	4.94	5.63	8.24	7.28	7.68	10.50	9.10	8.20	3.31	2.71	6.82	8.66	8.26
Na ₂ O	3.50	4.99	4.88	2.13	3.73	2.75	3.59	4.65	5.58	3.49	2.96	3.59	2.58
K ₂ O	1.58	0.26	0.17	4.25	0.40	0.85	0.15	0.34	1.08	1.01	1.24	0.61	2.57
P ₂ O ₅	0.21	0.19	0.05	0.15	0.14	0.12	0.12	0.17	0.13	0.04	0.16	0.20	0.12
LOI	7.52	7.62	6.57	8.34	3.88	4.21	4.63	3.82	3.77	3.61	7.42	7.94	3.38
#Mg	55	54	43	54	54	53	56	48	49	36	48	46	52
ppm													
Rb	53	11	10	99	15	24	10	8	36	28	66	34	77
Sr	109	130	50	103	211	209	129	137	216	35	170	134	307
Zr	74	71	98	66	62	53	50	78	56	148	83	63.8	44
Ba	294	102	38	488	105	167	51	126	414	242	119	72	313
Pb	-	-	-	-	-	-	-	1.99	-	-	5.73	3.22	-
Th	-	-	-	-	-	-	-	1.88	-	-	1.97	1.75	2.07
U	-	-	-	-	-	-	-	0.46	-	-	0.48	0.50	0.50
Nb	<10	<10	<10	<10	<10	1.99*	1.99*	3.5	2.74*	<10	1.82	2.66	1.00
Cr	189	188	56	103	199	165	184	126	163	14	109	85	90.2
Ni	63	58	10	29	62	53	63	26	34	14	35	20	18
Sc	36	36	10	32	40	35	34	38	39	11	36	30	36.6
V	230	254	42	203	267	274	261	259	240	27	259	259	294
Zn	145	107	10	65	80	76	188	91	107	99	86	81	115
Hf	-	-	-	-	-	-	-	1.63	-	-	1.77	1.68	1.38
Y	25	23	23	22	20	18	20	22	18	42	23	19.2	18
La	12	10	16	9.0	7.7	6.5	6.1	10.6	14	26	10	9.3	8.69
Ce	27	23	36	21	18	15	14	22.1	31	54	21.2	21.7	21.2
Pr	3.3	2.9	3.8	2.5	2.1	1.7	1.5	2.95	3.9	6.1	2.91	2.95	15.2
Nd	16	14	17	12	11	9.4	9.1	12.9	18	26	12.6	12.9	2.83
Sm	3.9	3.5	3.8	3.0	2.7	2.4	2.4	3.15	4.8	5.8	3.09	3.18	0.96
Eu	1.2	1.1	0.79	0.9	0.94	0.96	0.82	1.13	0.75	1.3	0.90	0.94	2.5
Gd	4.1	3.7	3.6	3.2	3.1	2.7	2.7	3.48	5.7	5.8	3.41	3.62	1.43
Tb	0.75	0.69	0.61	0.59	0.58	0.56	0.49	0.52	1.24	1.1	0.53	0.56	-
Dy	4.0	3.7	3.8	3.3	3.2	2.8	2.8	3.27	7.8	6.2	3.33	3.3	-
Ho	0.79	0.75	0.78	0.65	0.65	0.59	0.56	0.69	2.0	1.3	0.72	0.74	0.59
Er	2.4	2.2	2.4	2.0	2.0	1.7	1.7	1.89	5.6	4.0	1.95	1.97	-
Tm	0.50	0.50	0.50	0.50	0.50	0.50	0.50	0.28	0.87	0.62	0.30	0.3	-
Yb	2.3	2.1	2.5	1.9	1.9	1.7	1.6	1.8	5.4	4.0	1.92	1.88	-
Lu	0.50	0.50	0.50	0.50	0.50	0.50	0.50	0.27	0.76	0.61	0.28	0.28	0.265

* Sample/or element analyzed in the Instituto de Geociências, Universidade de São Paulo.

- No data.

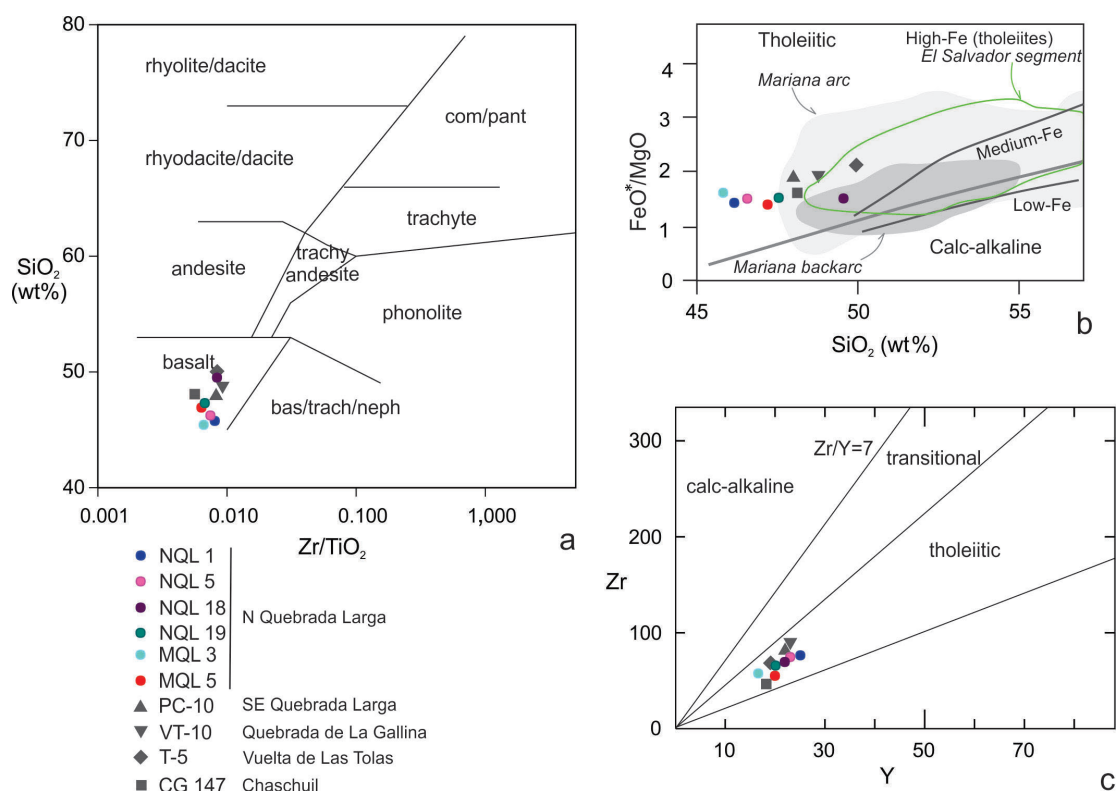


FIG. 11. Basalts and their related fragmented facies plotted in **a.** Zr/TiO₂ versus SiO₂ diagram (after Winchester and Floyd, 1977; Pearce, 1996); **b.** SiO₂ versus FeO*/MgO diagram (after Miyashiro, 1974). Mariana basin data from Ribeiro *et al.* (2013); El Salvador Segment (Central American Volcanic Arc) data from Patino *et al.* (2000); **c.** Y versus Zr diagram (after Barrett and McLean, 1999).

concentrations are also characteristics of the low-Ti series (cf. Cummins *et al.*, 1992; Peate *et al.*, 1992), as observed in abundances of Cr=85-199 ppm, Ni=18-63 ppm, Zr=44- 78 ppm, Nb<10 ppm, Hf=1.38-1.77 ppm, Th=1.75-2.07 ppm, U=0.46-0.50 ppm, Sr=103-211 ppm, Ba=51-488 ppm, V=203-274 ppm and Sc=30-40.

The chondrite-normalized REE patterns (after Boynton, 1984) of the analyzed rocks are shown in figure 12a. The ratio of Sm_N/Yb_N=0.29-0.41, also showing a slight enrichment of LREE relative to the HREE, with La_N/Yb_N=0.80-1.27. The moderate or absent Eu anomaly (Eu/Eu*=0.80-1.15) denotes a slight or no previous fractionation of plagioclase in the basalts.

In a spider diagram, trace elements normalized to N-MORB (after Sun and McDonough, 1989), the basalts (Fig. 12b) display features corresponding to

volcanic arcs basalts, such as Nb depletion respect of Ce and enrichment in Rb and Ba. The relatively high contents of Pb determined from some samples (Table 1) are typical of subduction environments. The tenors of Ti and Y are comparable to those of MORB, but also compatible with basalts from volcanic arcs (Pearce *et al.*, 1984). Moreover the contents of rare earths have an almost flat design (Fig. 12b), with slight slope increases towards the LREE with enrichments of 1 to 1.5 times MORB values.

The tectonic environment for the emplacement of the analyzed basalts can also be studied in different discrimination diagrams. In the V-Ti/1000 diagram (Shervais, 1982) the studied basalts plot between the fields of the arc tholeiitic basalt and the MORB-BAB ones (Fig. 13a) and in the Ti-Zr diagram (after Pearce, 1982; Condie, 1989) they

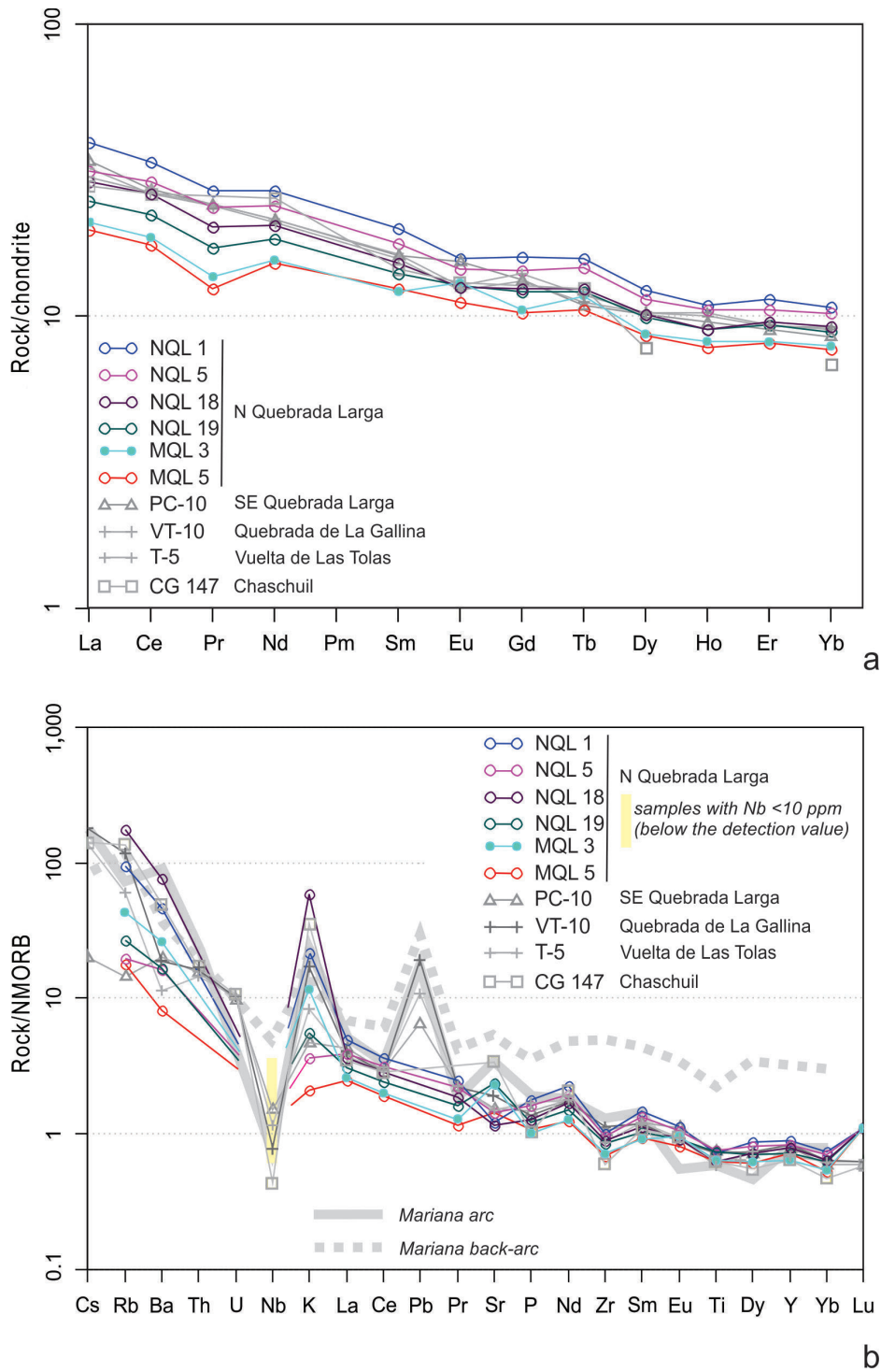


FIG. 12. Multi-element (spider) diagrams for the basalts and their related fragmented facies. **a.** Chondrite-normalized REE plots (after Boynton, 1984); **b.** N-MORB-normalized plots using normalizing factors of Sun and McDonough (1989). Mariana basin data from Ribeiro *et al.* (2013).

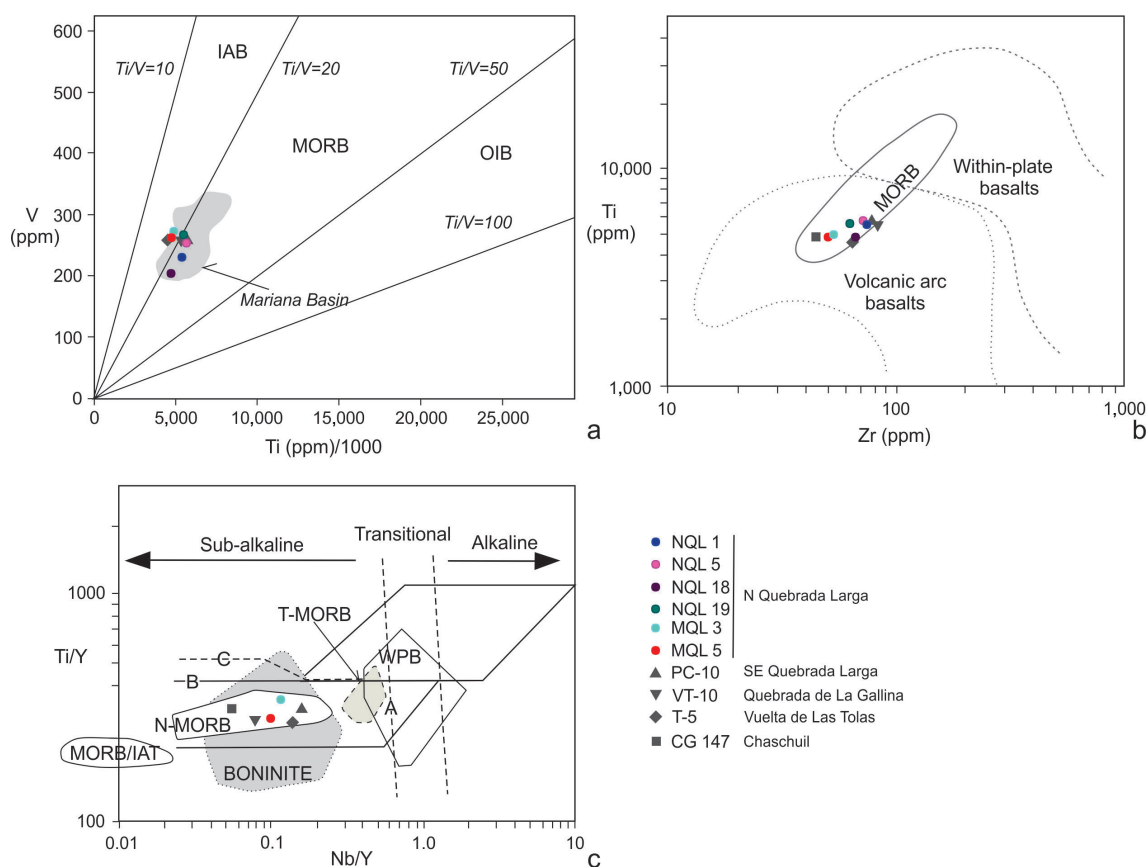


FIG. 13. **a.** Ti versus V diagram (Shervais, 1982). Mariana basin data from Ribeiro *et al.* (2013); **b.** Ti-Zr diagram (after Pearce, 1982; Condie, 1989) where the analyzed basalts plot along the MORB field; **c.** Ti/Y versus Nb/Y diagram (modified from Pearce, 1982) without samples with Nb < 10 (Table 1), to not produce erroneous interpretations. Fields: A: within plate; B: mid-ocean ridge; C: volcanic arc; MORB/IAT, mid-oceanic ridge basalt/island arc tholeiite; N-MORB, normal MORB; T-MORB, transitional MORB; WPB, within-plate basalts. Mariana basin data from Ribeiro *et al.* (2013).

fall in the MORB field within the volcanic arc basalts (Fig. 13b). In the Ti/Y-Nb/Y -diagram (Pearce, 1982) the samples plot along the N-MORB field (Fig. 13c).

5. Discussion

The basalts and the associated autoclastic to hyaloclastic facies studied in this paper integrate the lower portion of the Arenigian succession that crops out along the Las Planchadas range. They were studied together with other basaltic volcanic and volcanoclastic rock successions that crop out to the south (Fig. 2) and were compared with the Ordovician volcanic rocks located to the northwestern

part of the Puna to define their evolving tectonic setting at that time.

5.1. Emplacement environment

All of these Arenigian deposits are composed by facies associations of a subaqueous depositional environment, as shown by the well-preserved structures (lava flows, autoclastic facies and the related basaltic hyaloclastites).

Magma-water interaction can produce a wide variety of volcanic textures and structures ranging from pillow lavas, autobrecciation of lava flows and hyaloclastites formed by non-explosive interaction and fine-grained widely dispersed pyroclasts formed

by violent hydrovolcanic explosions (Carr and Jones, 2001). The structural and textural characteristics of the analyzed deposits can be generated by non-explosive magma-water interaction processes. The autoclastic basaltic facies, closely related to the massive lavas bodies, testify for a non-explosive fragmentation. Basaltic flows may be cooler and more viscous and may be subject to locally higher strain rates than other emplacements of massive basalts. They respond to stress in a brittle mode and develop blocks of lava, located at the margins of the lava bodies or dislodged by continuous movement of the flow (cf. McPhie *et al.*, 1993). In the study area, evidence for the cited process is the presence of autoclastic basaltic facies where the basaltic lavas are intensively fractured and show jigsaw-structures, with blocky clasts with curvilinear fracture surfaces and non-quenched margins.

The presence of Arenigian basaltic pillow breccias to the south of the study area (Quebrada Larga) also point to the non-explosive magma-water interaction.

To the south, along the Quebrada Larga area (Fig. 2), the same Arenigian succession is composed by basalts and their autoclastic facies, hyaloclastites and pillow breccias (Cisterna *et al.*, 2010a) in close spatial relationship to marine fossiliferous sediments (Fig. 3). It is possible to compare these deposits, based on their close spatial and temporal relationships, with ancient volcanic successions where basaltic massive and fragmented lava flows are dominant in the proximal section and replaced in the most distal part by pillow lavas and/or pillow fragmented breccias (Quebrada Larga-de La Gallina area) (cf. Dimroth *et al.*, 1978; Waters and Wallace, 1992).

5.2. Nature of the monomictic basaltic breccias

Hyaloclastites are composed by clastic aggregates formed by the non-explosive fracturing and disintegration of quenched lavas (cf. Rittmann, 1962). Autobrecciation and quench fragmentation probably operate together during the disintegration process of the subaqueous lava flow. Quenching produces fractures affecting the surface of the lavas. The clasts are formed *in situ* by the intersection of these fractures along the lava flow and by the spalling of quenched glass fragments (cf. McPhie *et al.*, 1993). These clasts may range widely in size from less than one millimeter to tens of centimeters

and may be entirely or partly glassy, vesicular or non-vesicular (cf. McPhie *et al.*, 1993).

The exposed basaltic breccias in the study area are monomictic with clasts shapes that vary between slabby, flow-foliated with jagged ends, ragged or blocky clasts with curvilinear surfaces and display jigsaw structure. The clasts develop quenched margins and the fragment size may vary from less than one mm (millimeter) to 50 cm (centimeters). They show different textures, can be flow foliated, have a vitric groundmass or be highly vesicular. These rocks may be defined as *in situ* hyaloclastites because they display almost all the characteristics indicated for this type of deposits, as clasts with jigsaw texture, indicating that they remain in place while being fragmented (cf. McPhie *et al.*, 1993).

The poorly sorted and quite chaotic nature of the breccia deposits suggests that they are proximal and formed next to its source. Furthermore, the occurrence of angular and fusiform blocks, with their glassy rinds frequently preserved, suggest that the breccia deposits were syn-eruptive, preferentially formed in the same space opened up in the lava flow, the site where temporary vents cause the hydro-magmatic activity (cf. Carr and Jones, 2001).

The mixture of fragments, either massive and angular or highly vesiculated and with glass rinds (Fig. 9), may be related to hydromagmatic episodes chiefly due to water-magma interaction, that relieved the pressure on the underlying melt, leading then to the sudden unmixing of the volatile component and to consequent magmatic fragmentation (cf. Dellino and La Volpe, 1995).

5.3. Chemical features, magma source and tectonic setting

The hyaloclastites show some chemical characteristics (e.g., low content Fe_2O_3 , MgO and Al_2O_3) due to the partial mobilization of the ferromagnesian components by alteration processes. They also have higher amounts of Zr (average 123 ppm and in basalts 64.5 ppm), that may be cause for a preferential enrichment in heavy accessory minerals (e.g., zircon) in the fine matrix. For this reason they are not included together with the basalts in the discrimination diagrams and not take into account in the tectonic interpretations.

Based on field relationships, petrography and chemical characteristics the studied basalts were

correlated with the basalts and basaltic hyaloclastites exposed to the south of the Las Planchadas range (Quebrada Larga, Quebrada de La Gallina, Chaschuil and Vuelta de Las Tolas) with a similar geological context (Fig. 2). So their chemical data (unpublished) together with those obtained in the study area are considered to define the tectonic setting at the north portion of the Famatina System during Early Paleozoic.

Among the main geochemical characteristics of the studied basalts are Zr, Ti and Y contents (*e.g.*, Figs. 11 and 12) typical of depleted mantle sources (*cf.* Pearce *et al.*, 1984) and enrichment in LILE components (*e.g.*, average of Rb content=148.5 ppm and Ba/La ratio of 20.42), reflecting a subduction related signature (*cf.* Pearce *et al.*, 1984). Such features may be compared with basalts emplaced along arc-backarc basins systems, where the lava compositions are indicative of mid-ocean ridge basalt (MORB)-like sources with high water contents (*cf.* Taylor and Martínez, 2003). A depleted mantle source for the magma is also consistent with the observed Nb/Yb ratio displayed for the studied basalts, 1.34 in average (*cf.* Elliott, 2003; Pearce *et al.*, 2005).

The relatively low Nb/U ratio (7.61-2.00) compared to the oceanic basalts (MORB and OIB) with 47 ± 10 (Hofman *et al.*, 1986), may indicate enrichment of U relative to Nb by flux of subduction related fluids and melts into the depleted mantle (*c.f.* Pearce, 2008). The same is pointing by the Pb/U (4.33-11.94) and Pb/Ce (0.09-0.27) ratios (*c.f.* Pearce, 2008). In subduction related systems high degrees of melting are associated with low ratios of more incompatible to less incompatible elements of relatively low fluid mobility (*e.g.*, La/Yb, Sm/Lu, Nb/Y and Nb/Zr) (*c.f.* Carr *et al.*, 1990, 2003). For the studied basalts these ratios range between low to intermediate (average La/Yb=4.88, Sm/Lu=6.38, Nb/Y=0.11, Nb/Zr=0.03) pointing to variable degrees of melting. Relatively low values for the Sm/Lu ratio (3.62-11.67), U/Th (0.24-0.28) and Sr/Ce (4.04-14.48) in the analyzed basalts reflect fluid fluxes into the mantle wedge possibly resulting in higher degrees of melting (*e.g.*, Wehrmann *et al.*, 2014).

The analyzed basalts show a positive correlation between Th/Yb - Nb/Yb ratios (Fig. 14a) and fall on the volcanic arc domain with signature of a relatively depleted mantle source. On the TiO_2/Yb versus Nb/Yb diagram (Fig. 14b), they fall along the MORB field denoting relatively shallow melting and normal MORB-like source components (*cf.* Pearce, 2008).

This ambivalence is characteristic for subduction related tholeiitic basalts. Similar characteristics have been observed, for example in volcanic rocks of the Central America Volcanic Arc (Figs. 14a, b), in particular in the El Salvador-Guatemala segments. There, a transition from volcanic arc to an extension related volcanism (back-arc) along an overriding plate progressively thicker has been mentioned (Morgan *et al.*, 2008; Wehrmann *et al.*, 2014). Similar variation has also been observed in volcanic rocks of the Marianas volcanic-arc (Morgan *et al.*, 2008).

Marine sediments commonly have enrichment of elements that become fluid-mobile during subduction and their variations may be observed using ratios between highly fluid-mobile and less fluid-mobile incompatible trace elements. The relatively low values for the Ba/La (7.7-54.2), Ba/Nb (25.63-483.92) and U/Th (0.24-0.28) ratios of the basalts (Table 2) are comparable to the hemipelagic sediment from younger subduction systems such as the Central American Volcanic Arc (data from Heydolph *et al.*, 2012; Wehrmann *et al.*, 2014). The relatively high contents of Th (1.75-2.07 ppm), U (0.46-0.50 ppm) and Nb (1.00-3.05 ppm) in the basalts at the Las Planchadas range are supporting this idea.

All the previous geochemical considerations suggest for the studied basalts a depleted mid-ocean ridge basalt (MORB)-like source, modified by the contribution of subducted components derived from water rich marine sediments, features that are compatible with basalts emplacement along a back-arc basin.

As the back-arc-basin lavas (BAB) and mid-ocean ridge basalts (MORB) are considered to be produced by adiabatic decompression of depleted mantle (*cf.* Ribeiro *et al.*, 2013) and the arc and back-arc lavas formed by melting of a more hydrated and depleted mantle source (Reagan *et al.*, 2010), the studied basalts could be explained by a magma generated by an adiabatic decompression melting of an hydrated and depleted mantle (*cf.* Ribeiro *et al.*, 2013).

5.4. The Famatina System in a regional geologic framework

On the basis of the geochemical data here presented, together with the facies association observed in the Arenigian volcanic-volcaniclastic successions, it is possibly to consider the developing of a volcanic arc-back arc basin system during the Arenigian along

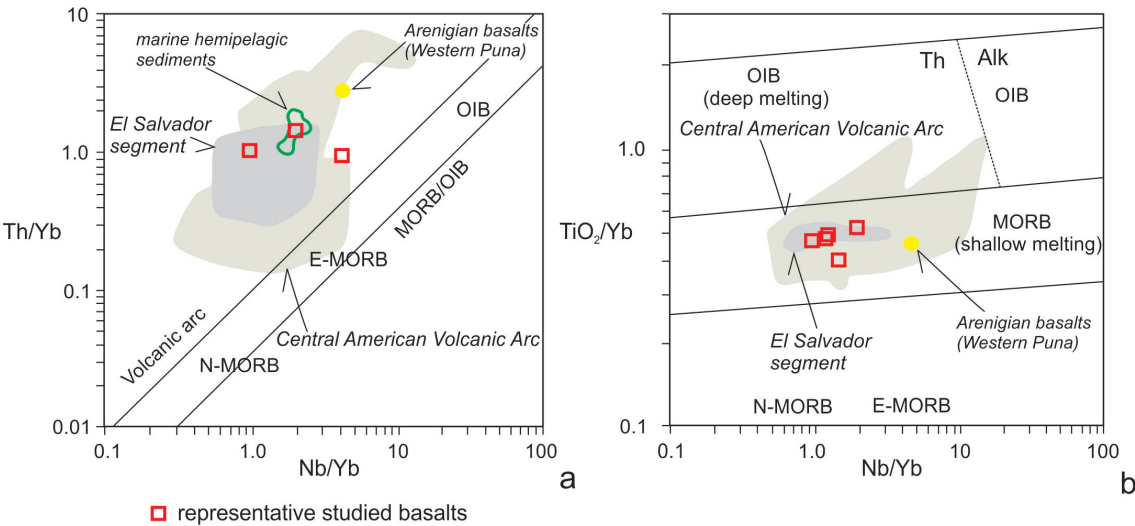


FIG. 14. **a.** Th/Yb versus Nb/Yb diagram (after Pearce, 2008). The compositions of input marine sediments are from Heydolph *et al.* (2012); **b.** TiO₂/Yb versus Nb/Yb diagram (after Pearce, 2008). Basalts from Western Puna data from Coira *et al.* (1999, 2009); El Salvador Segment and Central American Volcanic Arc data from Wehrmann *et al.* (2014).

TABLE 2. REE ELEMENTS RATIOS OF THE STUDIED BASALTS.

Location	Northern Quebrada Larga						Southeastern Quebrada Larga	Quebrada de La Gallina	Vuelta de Las Tolas	Chaschuil
Sample	NQL1	NQL5	NQL18	NQL19	MQL3	MQL5	PC-10*	VT-10*	T-5*	CG-147*
Rock type	basalt	basalt	basalt	basalt	basalt	basalt	basalt	basalt	basalt	basalt
Ba/La	24.5	10.2	54.2	13.6	25.7	8.5	11.9	11.9	7.7	36
Ba/Yb	127.80	48.57	256.84	55.26	98.23	31.87	70.00	61.98	38.30	-
Ba/Nb	-	-	-	-	483.92	25.63	36.00	65.38	27.07	313.00
La/Nb	-	-	-	-	0.31	0.33	0.33	0.18	0.29	0.11
La/Yb	5.22	4.76	4.74	4.05	3.82	3.81	5.89	5.21	4.95	-
La/Sm	3.08	2.86	3.00	2.85	2.71	2.54	3.36	3.24	2.92	9.05
Nb/Yb	-	-	-	-	1.17	1.24	1.94	0.95	1.41	-
Nb/Zr	-	-	-	-	0.04	0.04	0.04	0.02	0.04	0.02
Nb/U	-	-	-	-	-	-	7.61	3.79	5.32	2.00
Pb/Ce	-	-	-	-	-	-	0.09	0.27	0.15	-
Pb/U	-	-	-	-	-	-	4.33	11.94	6.44	-
Sm/Lu	7.80	7.00	6.00	5.40	4.80	4.80	11.67	11.03	1.69	3.62
Sr/Ce	4.04	5.65	4.90	11.72	13.93	9.21	6.20	8.02	6.17	14.48
Sr/Yb	47.39	61.90	54.21	111.05	122.94	80.62	76.11	88.54	71.28	-
Th/Nb	-	-	-	-	-	-	0.54	1.08	0.66	2.07
Th/Yb	-	-	-	-	-	-	1.04	1.03	0.93	-
TiO ₂ /Yb	0.40	0.46	0.41	0.49	0.48	0.49	0.53	0.47	0.40	-
Yb/La	0.19	0.21	0.21	0.25	0.26	0.26	0.17	0.19	0.20	-

* Sample/or element analyzed in the Instituto de Geociências, Universidade de São Paulo.

- No data.

the northern part of the Famatina System. Previously, during the Tremadocian, the volcanism in that region was related to the evolution of a marginal basin where depleted mantle derived melts (MORB) were produced and enriched through an extensional regime above the eastward dipping subducting slab (Cisterna *et al.*, 2010b). Similar characteristics have been observed for the Tremadocian volcanism occurred along the western slope of the Puna (Coira *et al.*, 2009).

During the Arenigian the magmatism at the western of the Puna was related to a volcanic arc-back-arc basin system, which evolved on attenuated continental crust (Coira *et al.*, 2009), environmental conditions similar to those proposed for the studied basalts at Las Planchadas range. Taking into account the ages, petrographic and chemical similarities between the Ordovician volcanic successions cropping out along the northern part of the Famatina System and the western magmatic Puna Belt (Fig. 1c) it is possible to argue a common tectonic evolution and to support the interpretations that propose the development of an Ordovician Puna-Famatina subduction related magmatic arc-back arc system along the western margin of Gondwana.

6. Conclusions

The volcanic sequence analyzed at the middle portion Las Planchadas range, northern part of the Famatina System, consists of basaltic lava flows, autoclastic breccias and hyaloclastites facies denoting a subaqueous depositional environment, as shown by the well-preserved structures indicative of non-explosive magma-water interaction. Sedimentary rocks, bearing Arenigian fossil fauna are intercalated in the upper part of the sequence.

The geochemical characteristics of the study basaltic rocks, such as LREE enriched REE patterns, enrichment in LILE, in particular K and Pb, and MgO, Fe₂O₃, Ti, Zr, Y, Nb contents indicate a depleted mid-ocean ridge basalt (MORB)-like source, modified by the contribution of subducted components derived from water rich marine sediments, features that are compatible with basalts emplacement along a back-arc basin.

Taking into account the ages, petrographic and chemical similarities between the Ordovician volcanic successions cropping out along the northern part of the Famatina System and the western magmatic Puna Belt it is possible to argue a common tectonic

evolution and to support the interpretations suggesting the development of an Ordovician Puna-Famatina subduction related magmatic arc-back-arc system along the western margin of Gondwana.

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